



**Juan R. Pardo**

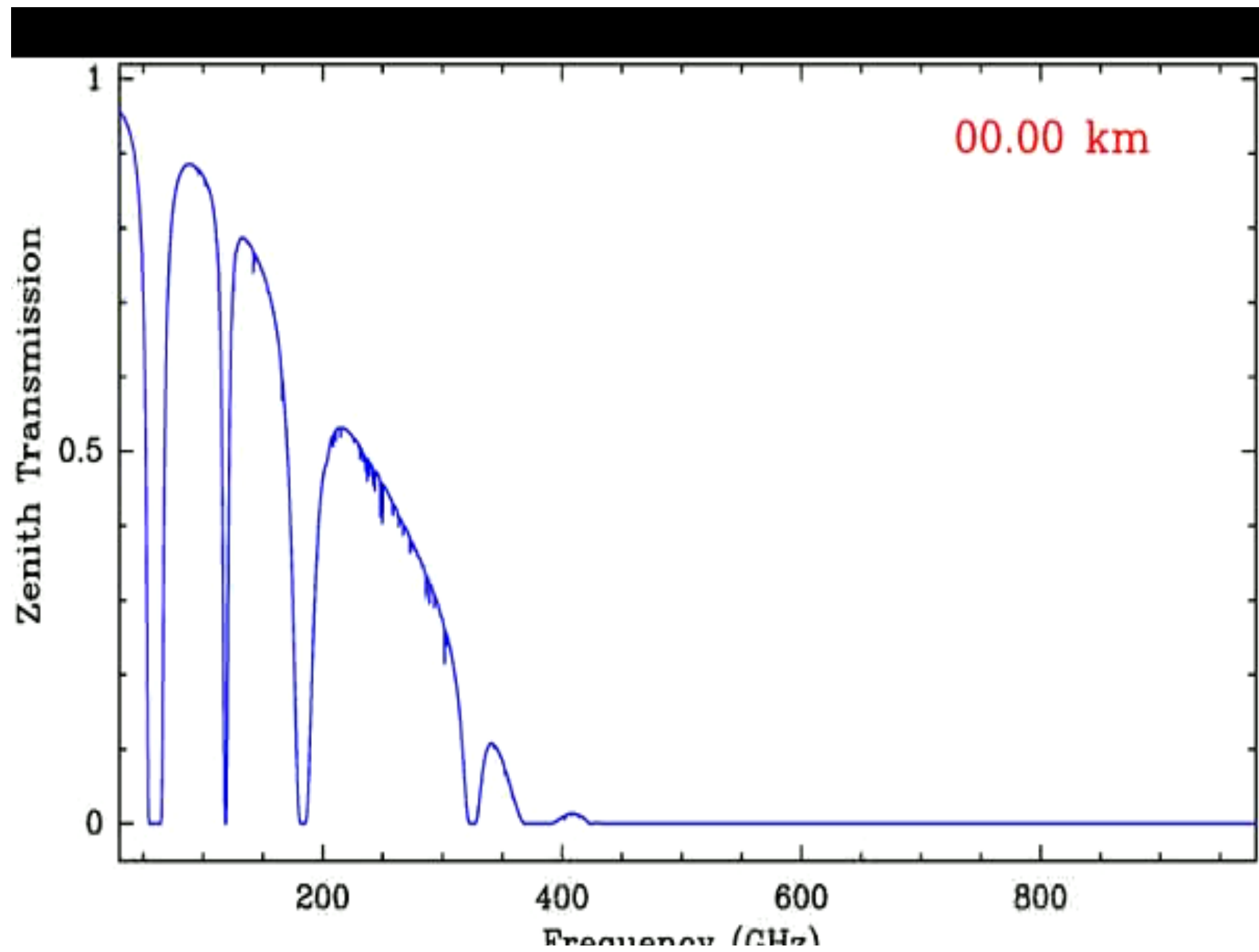
Consejo superior de Investigaciones  
Científicas (Spain)

**COSPAR Workshop, Quito**

**Mar 8th, 2018**

## **Atmospheric mm/submm refraction index**

- a. Physical basis**
- b. Experimental basis**
- c. Up-to-date model**
- d. Applications**



# General radiative transfer equation (using coordinates: $z$ , $\mu=\cos(\theta)$ , $\varphi$ )

$$\mu \frac{dI(z, \mu, \varphi)}{dz} = K(z, \mu, \varphi)I(z, \mu, \varphi) - \int_{-1}^1 d\mu' \int_0^{2\pi} d\varphi' Z(z, \mu, \varphi, \mu', \varphi') I(z, \mu', \varphi') - \sigma(z, \mu, \varphi) B[T(z)]$$

**I**=(**I**,**Q**,**U**,**V**)<sup>T</sup> Radiation field (Stokes column vector)

**K** 4x4 extinction matrix

**Z** 4x4 phase matrix (describing scattering)

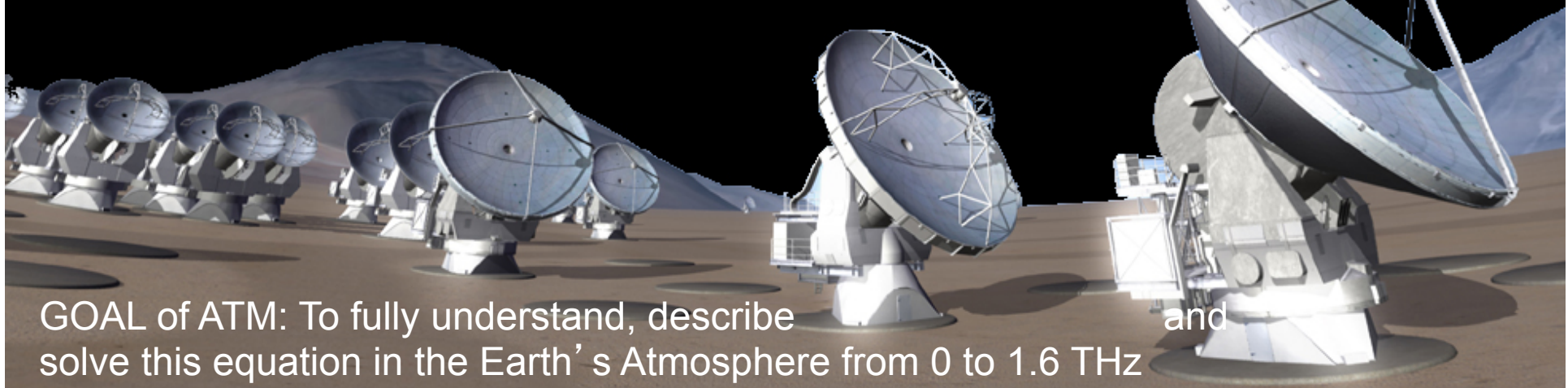
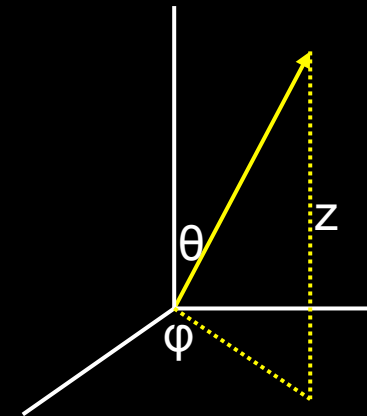
**σ** 4x1 emission column vector

**B** Blackbody radiance at temperature  $T$  (due to LTE)

Frequency dependence is implicit

$$K_{i1}(z, \mu, \varphi) = \int_{-1}^1 d\mu' \int_0^{2\pi} d\varphi' Z_{i1}(z, \mu, \varphi, \mu', \varphi') + \sigma_i(z, \mu, \varphi), i = 1, \dots, 4$$

Detailed energy balance



GOAL of ATM: To fully understand, describe and solve this equation in the Earth's Atmosphere from 0 to 1.6 THz

# Clear Atmosphere

$$\frac{dI_\nu(s')}{d\tau_\nu} = -I_\nu(s') + B_\nu(T[s'])$$

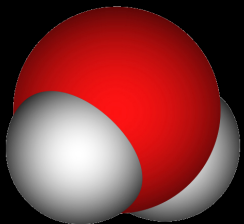
gas-phase  $\kappa_\nu$

$s'$  coordinate along the path;  
 $S_\nu = \epsilon_\nu / \kappa_\nu$  source function;  
 $d\tau_\nu = \kappa_\nu ds$  differential opacity.

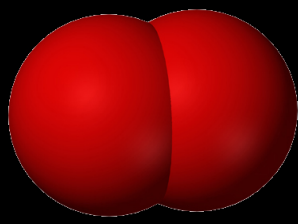
$$(\kappa_\nu)_{lu} = \frac{8\pi^3 N_\nu}{3hcQ} \left( e^{-E_l/kT} - e^{-E_u/kT} \right) \cdot |\langle u | \mu | l \rangle|^2 f(\nu, \nu_{l \rightarrow u})$$

Rotational lines  
absorption

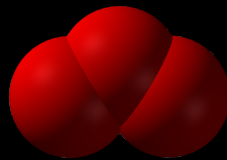
Continuum-like  
absorption



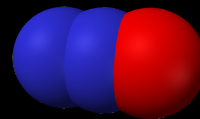
H<sub>2</sub>O



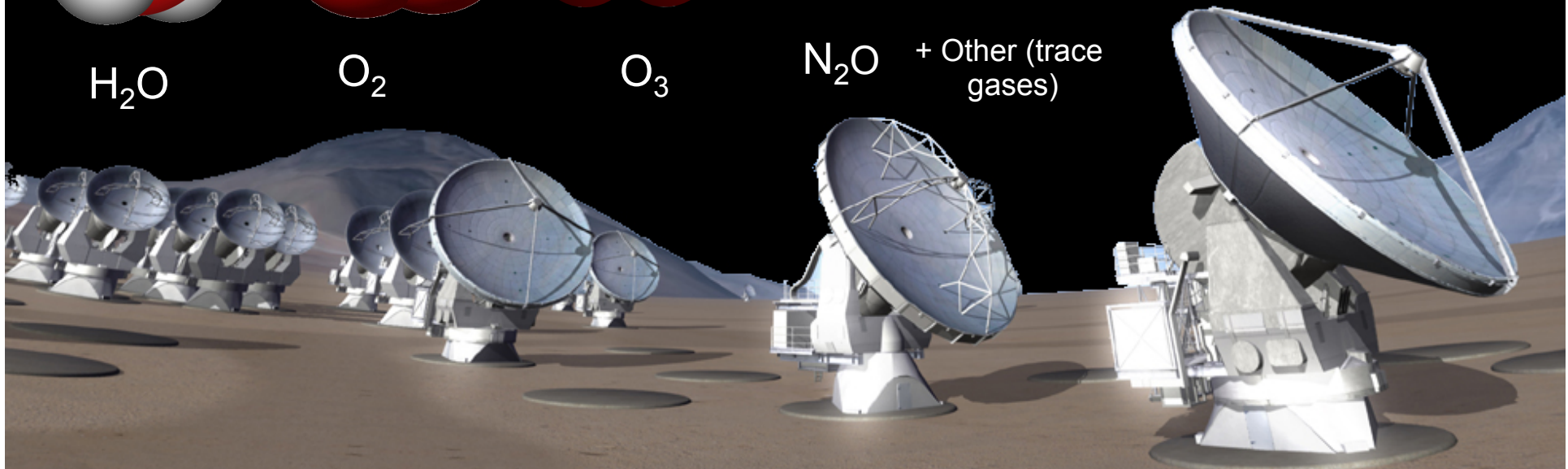
O<sub>2</sub>



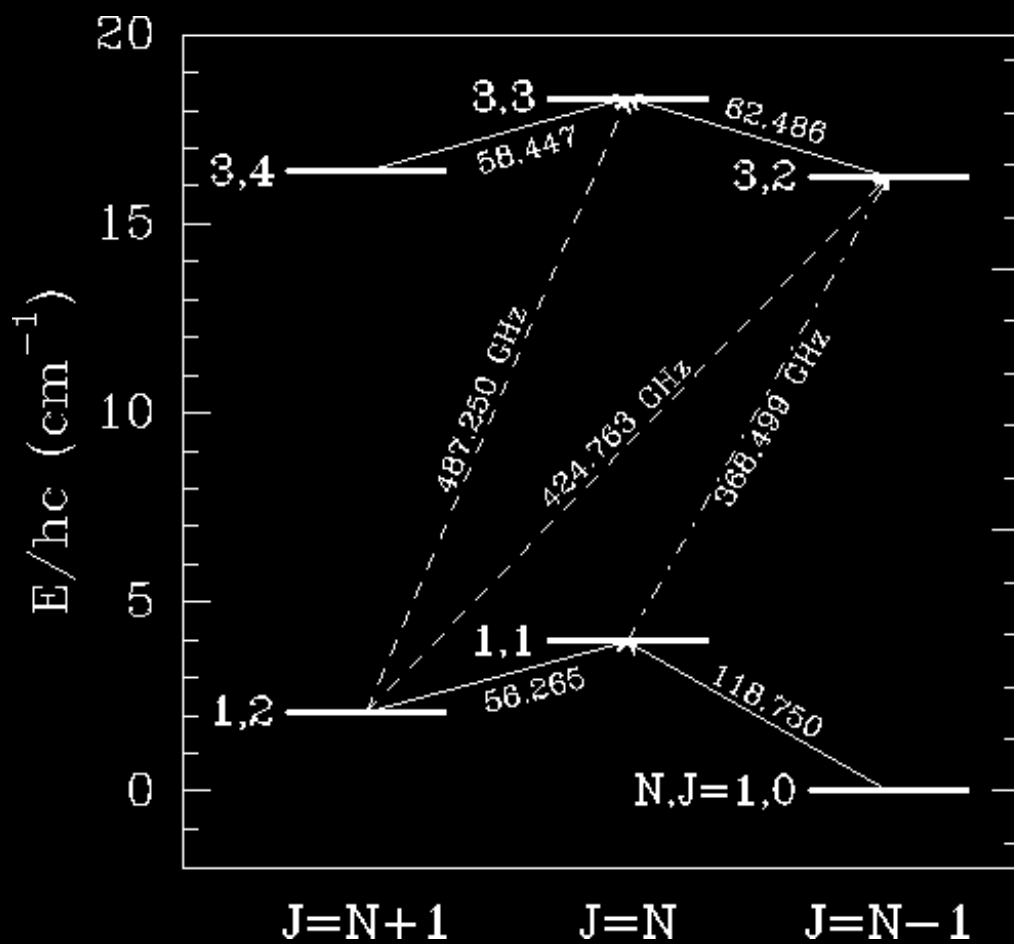
O<sub>3</sub>



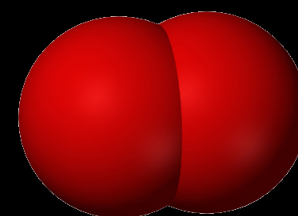
N<sub>2</sub>O + Other (trace  
gases)



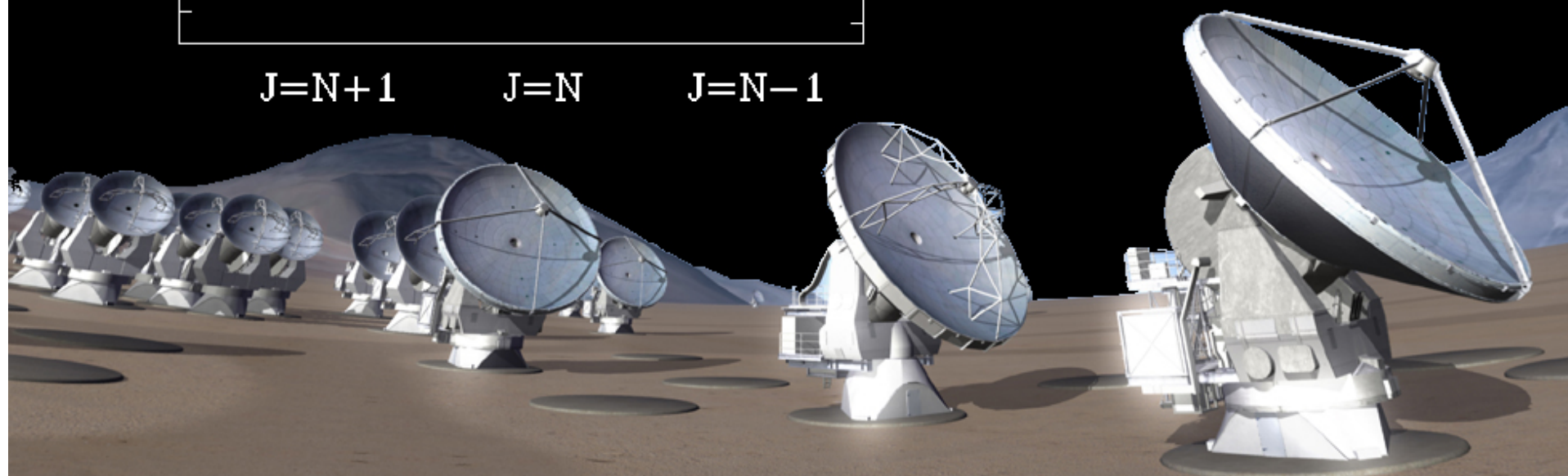


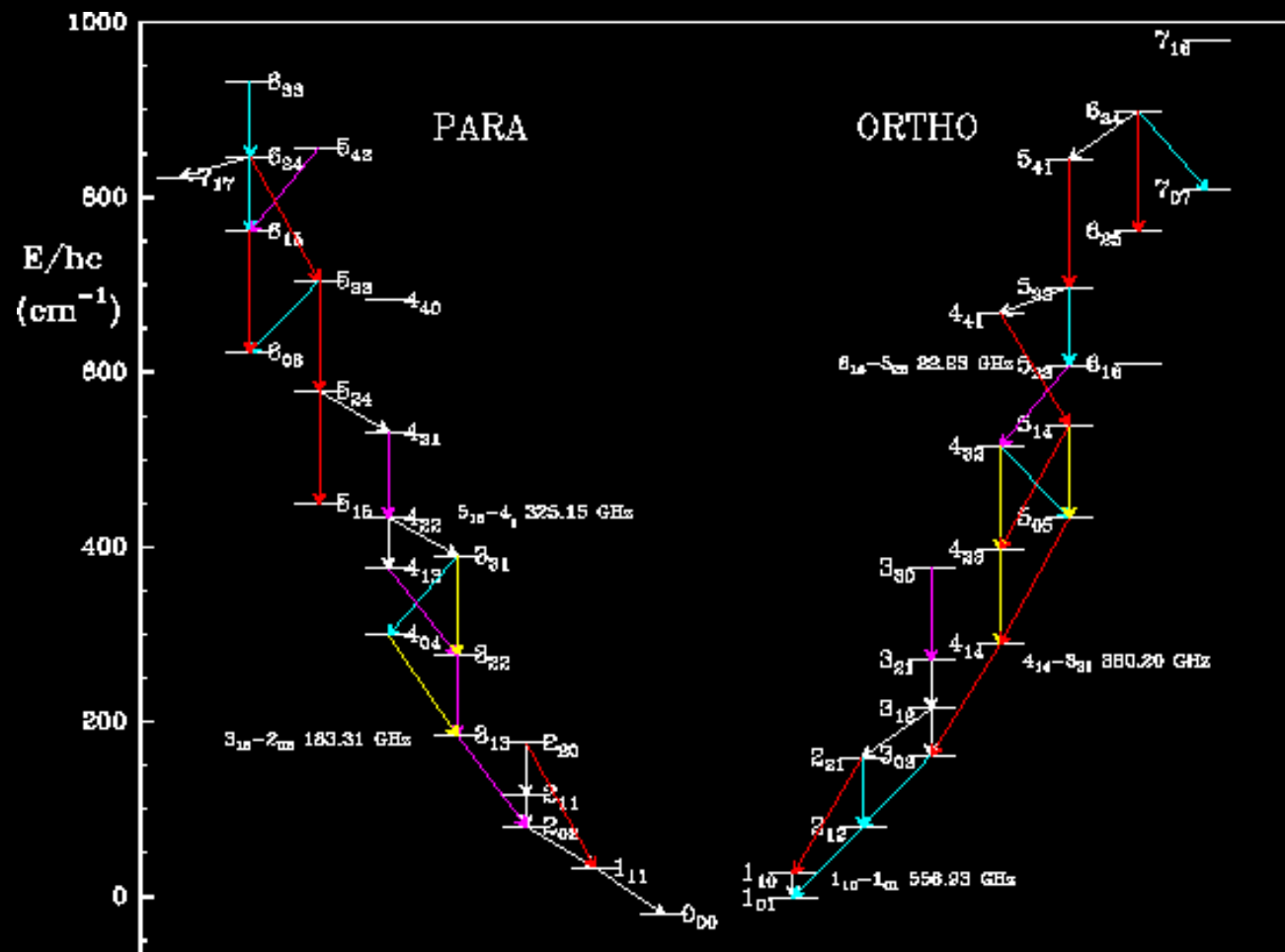


Lowest  
rotational  
transitions  
(M1)

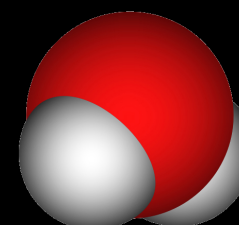


$O_2$

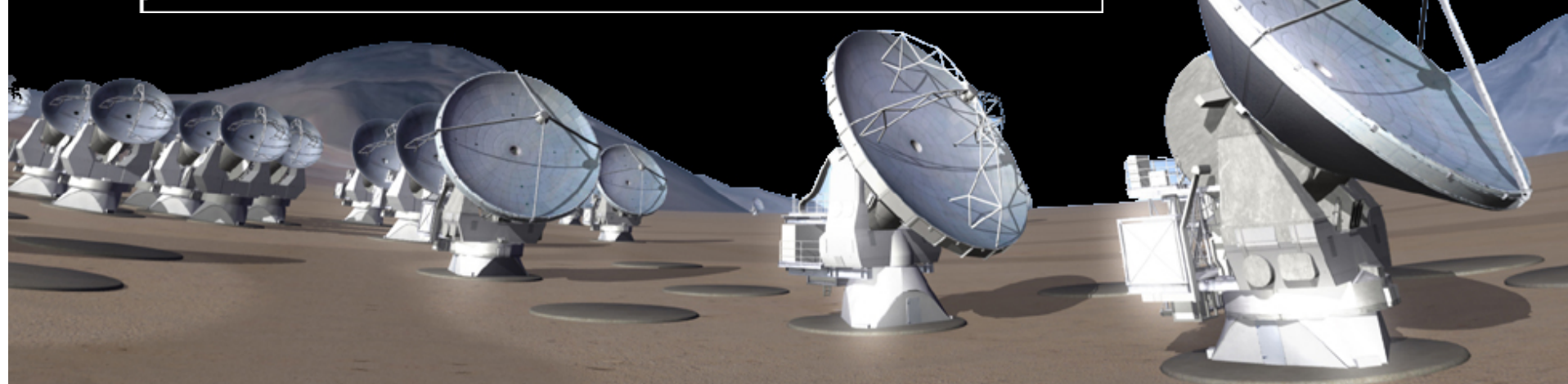




rotational  
transitions  
(E1)



$\text{H}_2\text{O}$



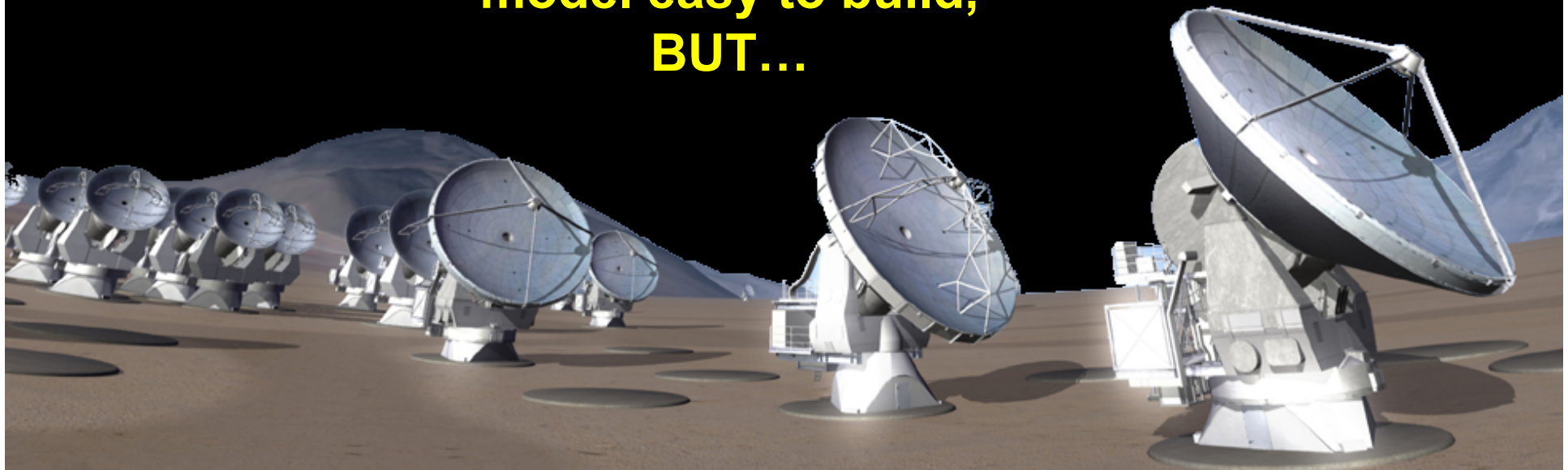
**Line strengths  
from quantum  
mechanics**

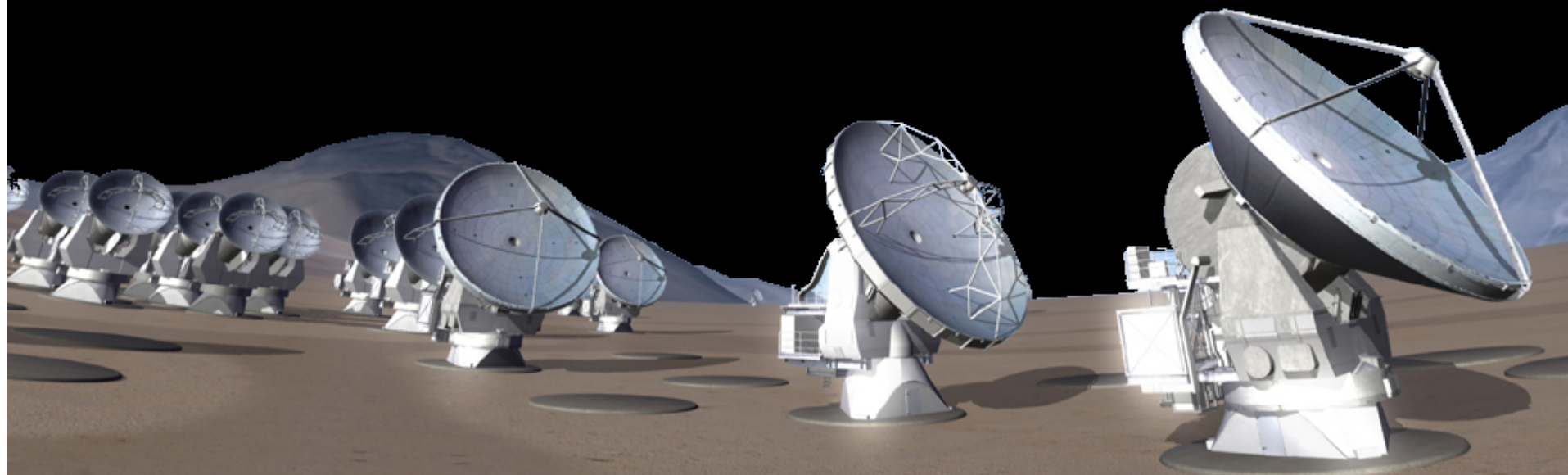
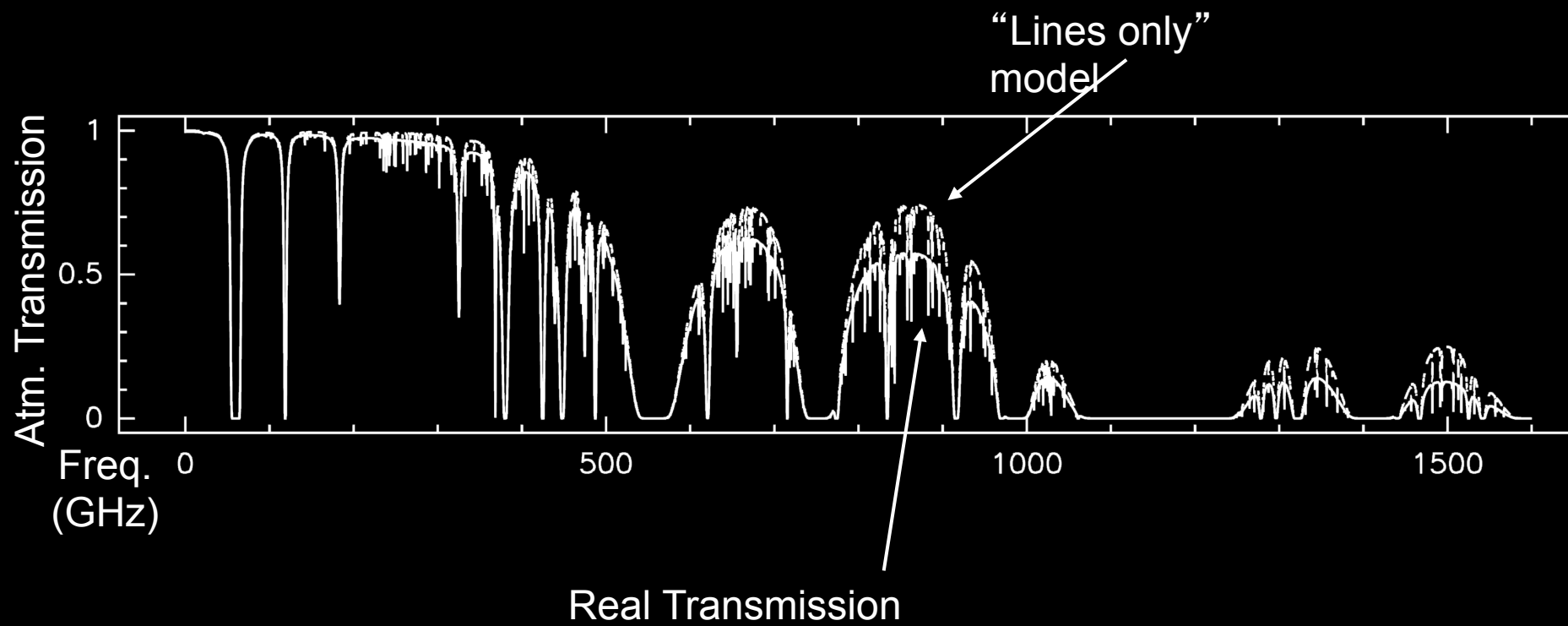
**+**

**Line profiles (line  
widths from laboratory)**



**Simple transmission  
model easy to build,  
BUT...**



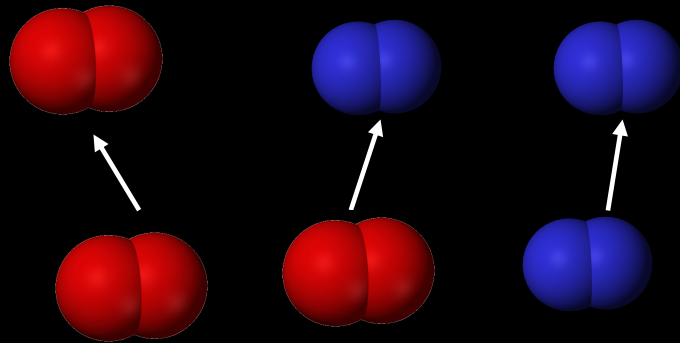




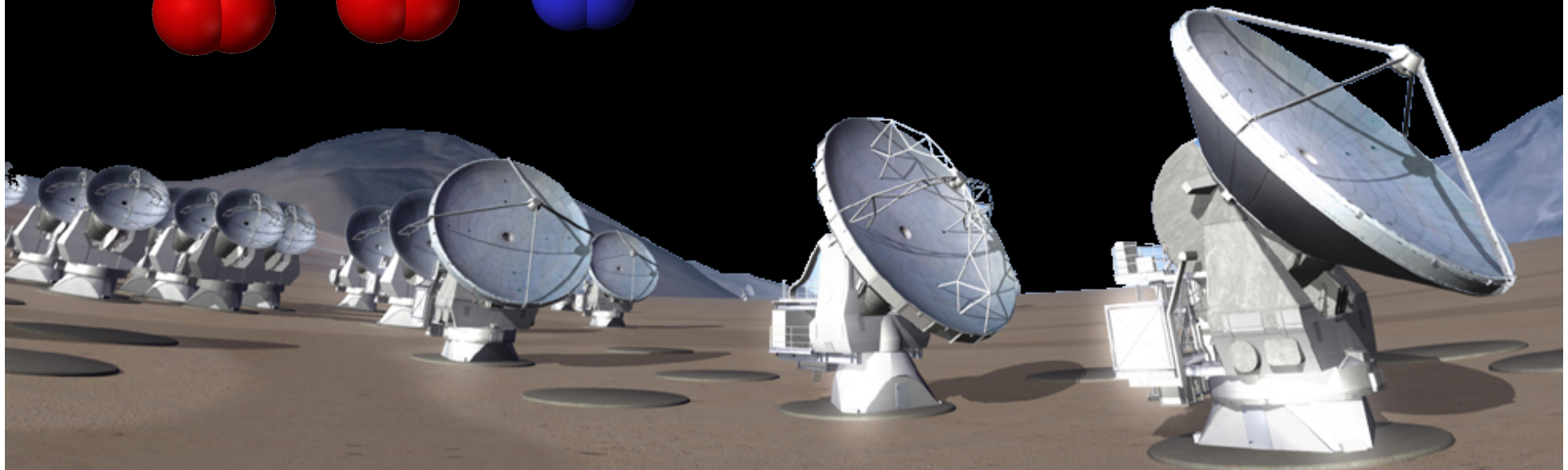
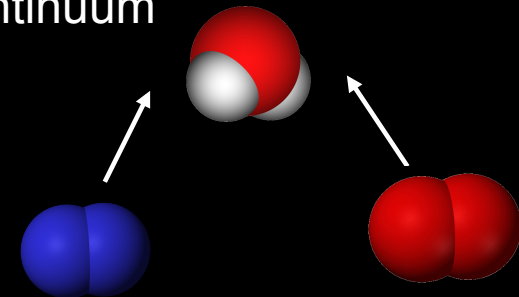
Today we know that the nature of “continuum-like” absorption in clear sky conditions is collision induced absorption involving pairs such as  $\text{O}_2\text{-O}_2$ ,  $\text{N}_2\text{-N}_2$ ,  $\text{O}_2\text{-N}_2$ ,  $\text{O}_2\text{-H}_2\text{O}$  and  $\text{N}_2\text{-H}_2\text{O}$

## Collision-induced absorption

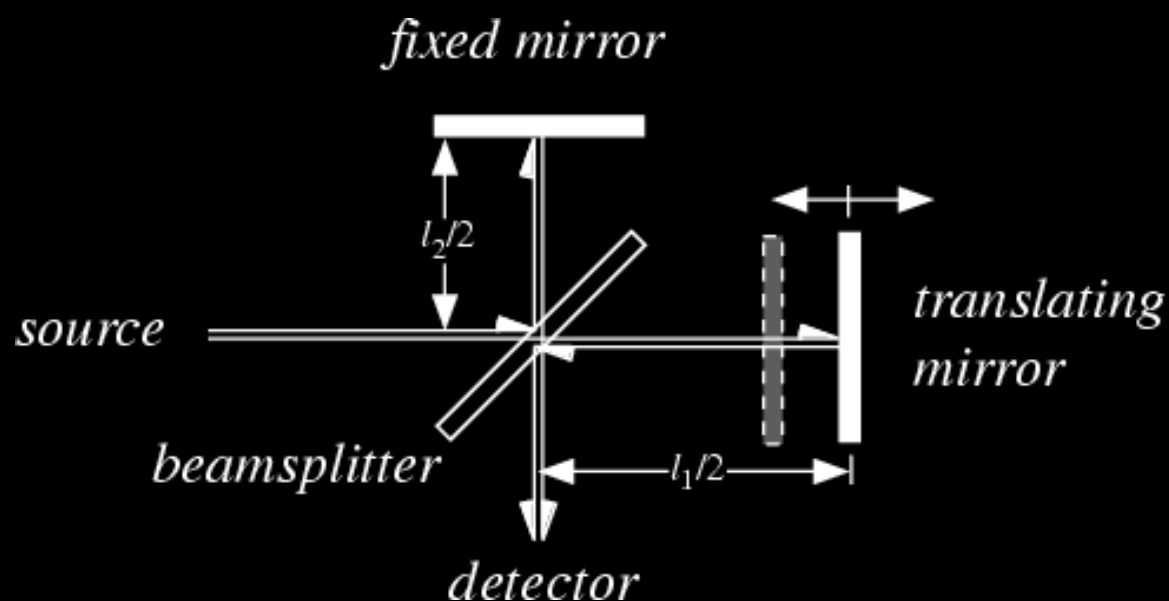
“Dry”  
continuum



“Wet”  
continuum

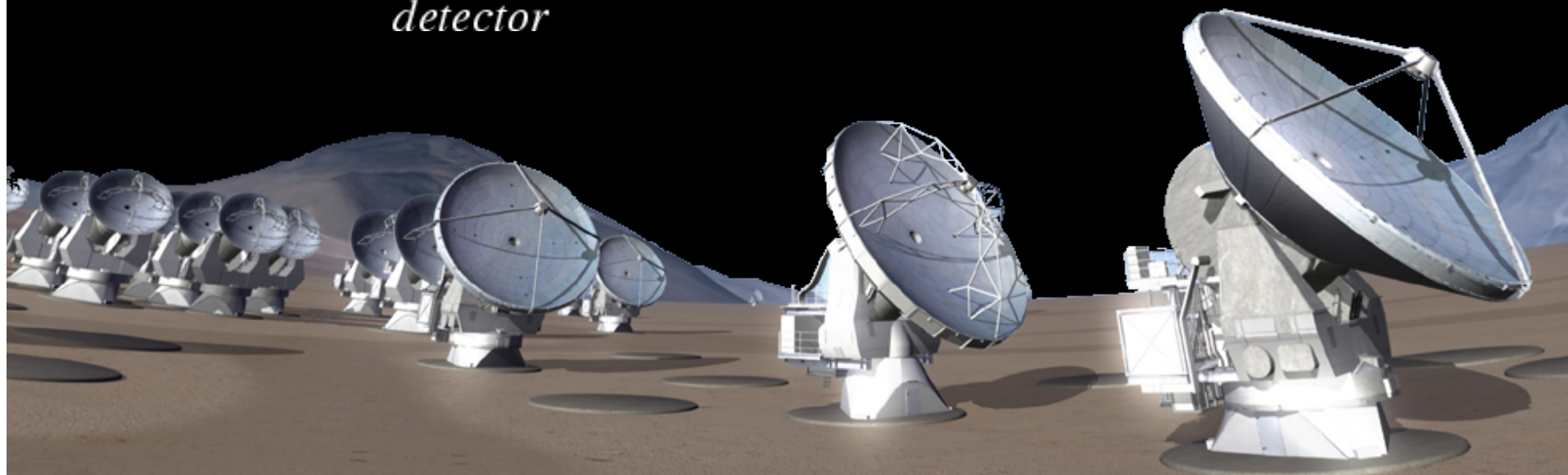


## Experimental basis for refined models: Direct measurements with FTS experiments at Mauna Kea, Chajnantor & Sout Pole



### Characteristics of CSO-FTS

- Mounted on Cassegrain focus of telescope for dedicated obs. runs.
- Detector:  $^3\text{He}$  cooled Bolometer
- Moving arm: 50 cm  
~ 200 MHz resolution
- Filters: 7 different (165 to 1600 GHz)





**Mauna Kea (4200 m, -5 °C), Hawai'i**



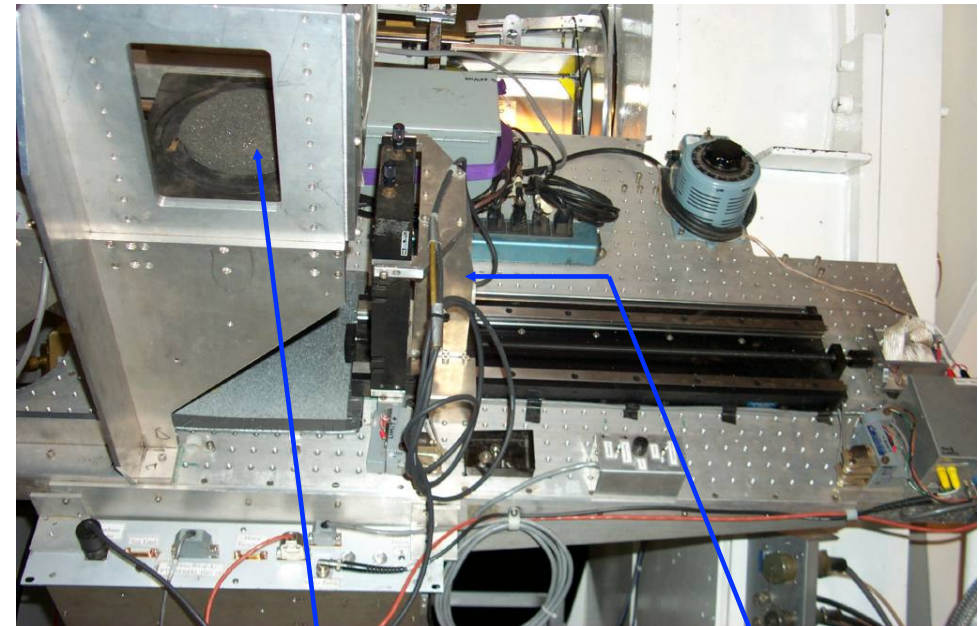
**Hapuna Beach (0 m, 25 °C, Hawai'i)**



# Caltech Submillimeter Observatory - Fourier Transform Spectrometer (CSO-FTS)



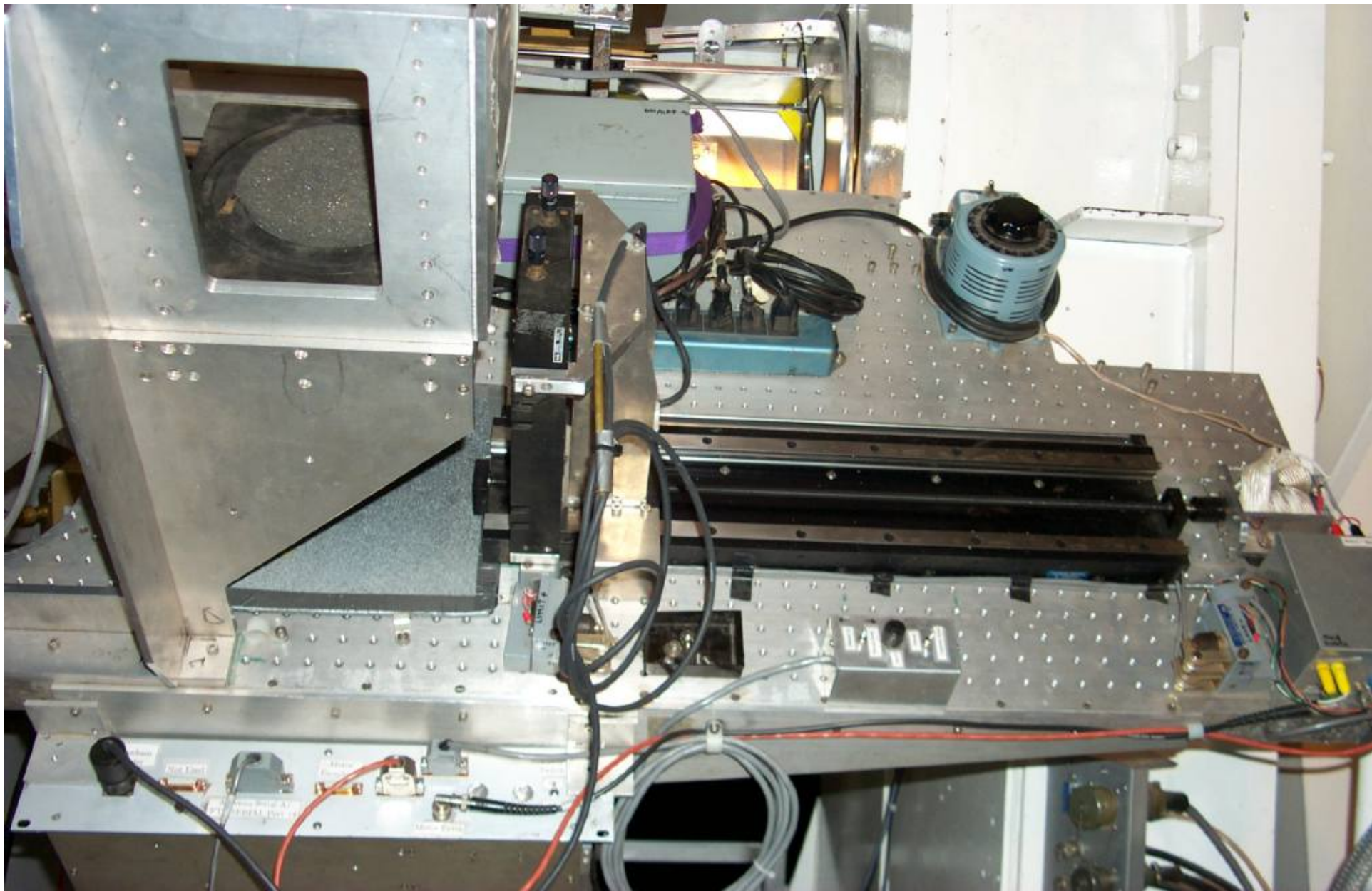
Last mirror and bolometer (cooled to liq.  $^3\text{He}$ )



Fixed mirror (can rotate for holography)

Moving mirror





# **CSO-FTS approach to solve the « excess of continuum » problem**

**Use well calibrated measurements acquired  
during very dry conditions**

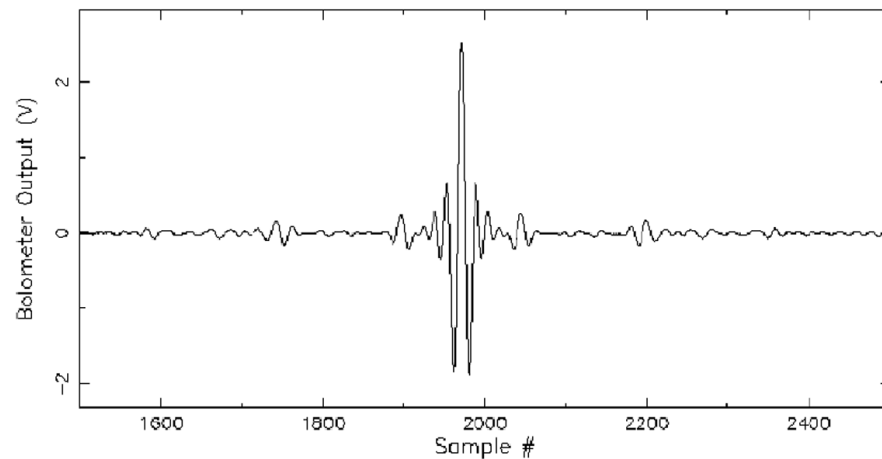
- $T_0 \sim 270 \pm 3 \text{ K}$
- $P \sim 620 \pm 1.5 \text{ mb}$  at Mauna Kea summit.

**As a consequence:** The “dry” atmospheric absorption is basically the same (within 1-3 %) in the different situations.  
The remaining opacity is proportional to the PWV

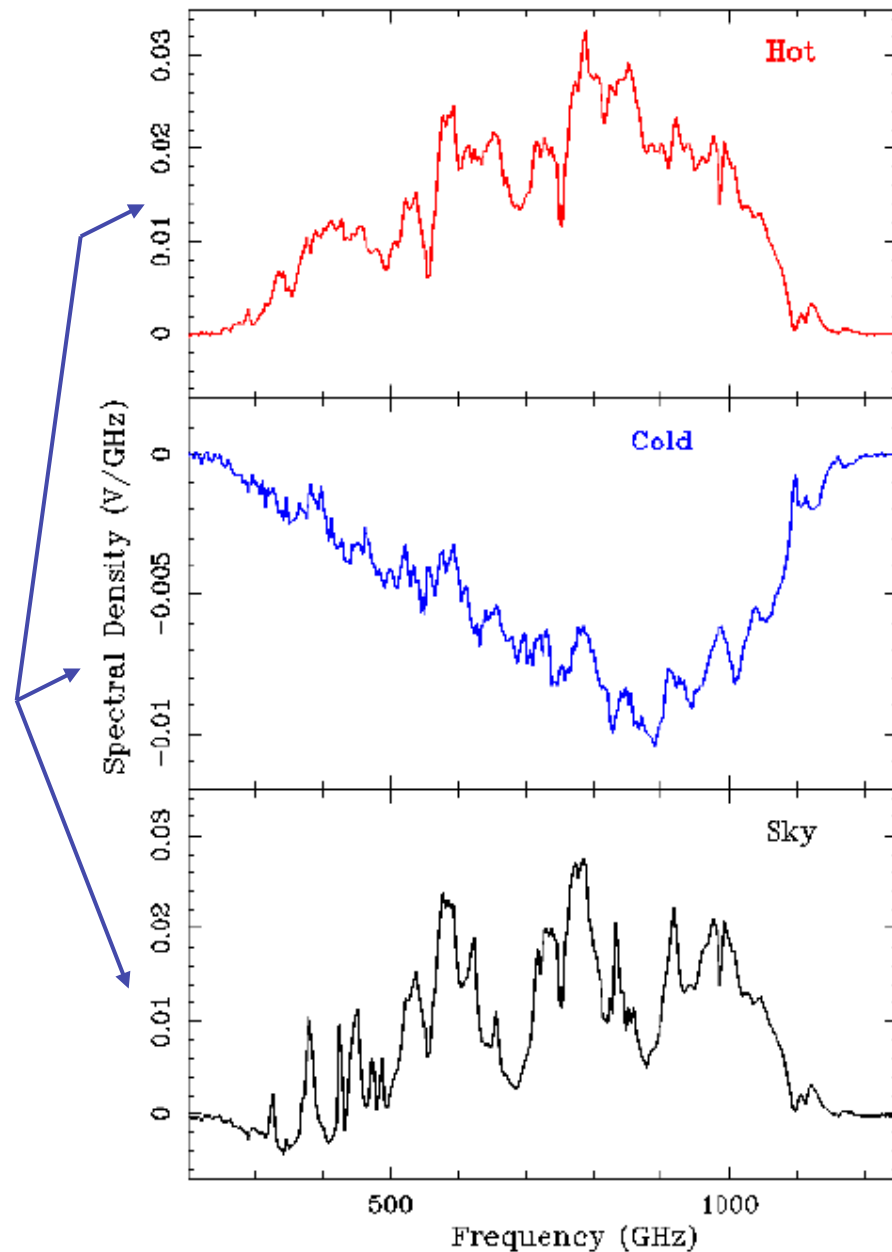
**The water vapor column can be determined very precisely from the near wings  
of water lines, virtually independently of the continuum terms.**

# CSO-FTS. Calibration of atmospheric measurements

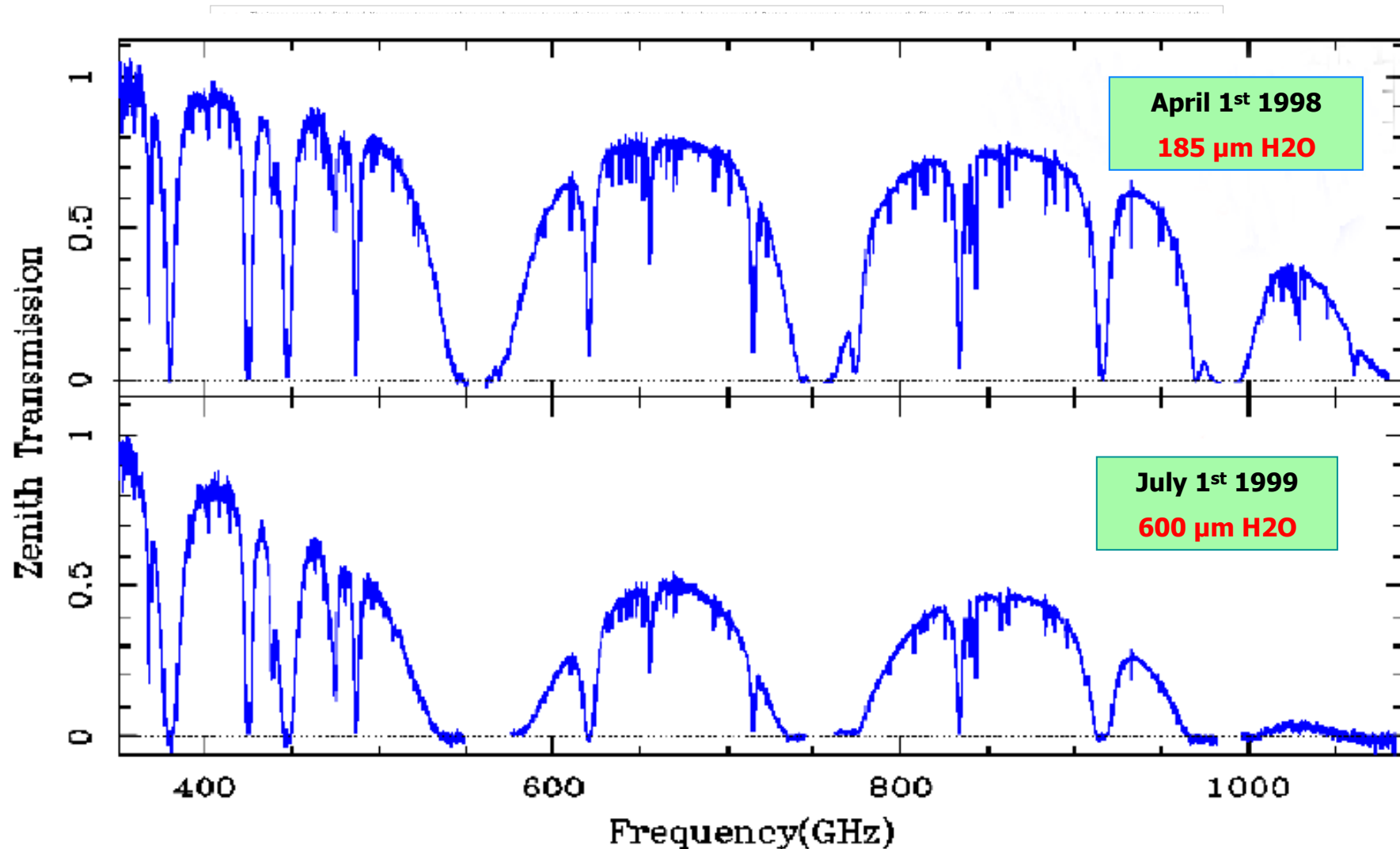
**An accurate calibration of the spectra is key to the atmospheric goals of the experiment**



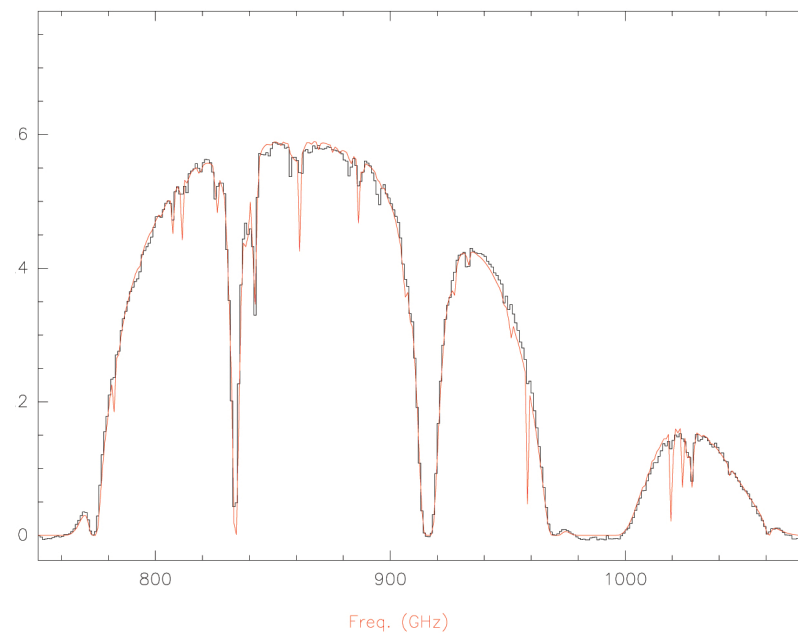
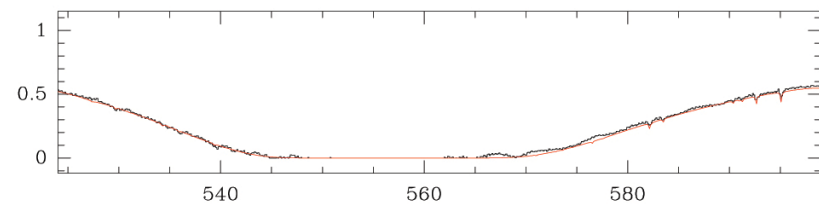
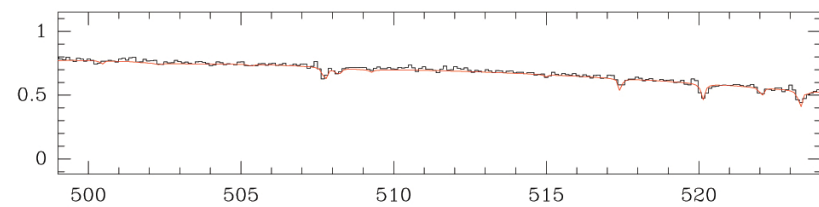
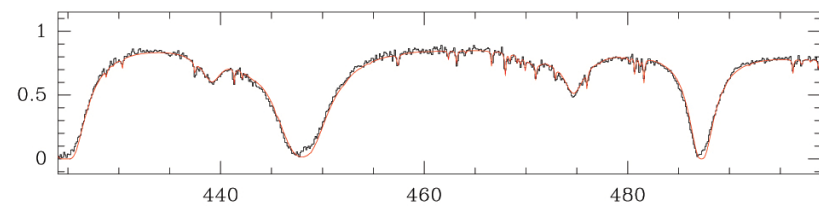
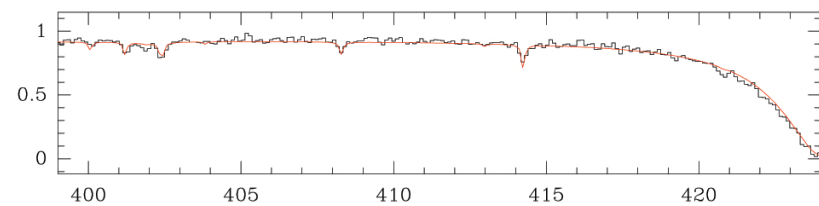
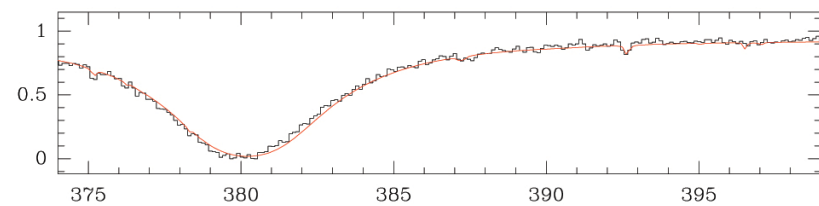
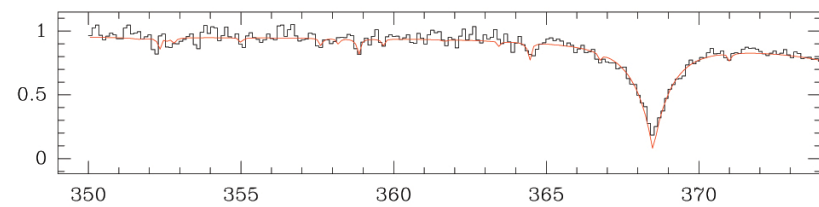
**Raw data of CSO- FTS experiment  
(full resolution two-sided scan  
takes about 5 minutes to complete)**

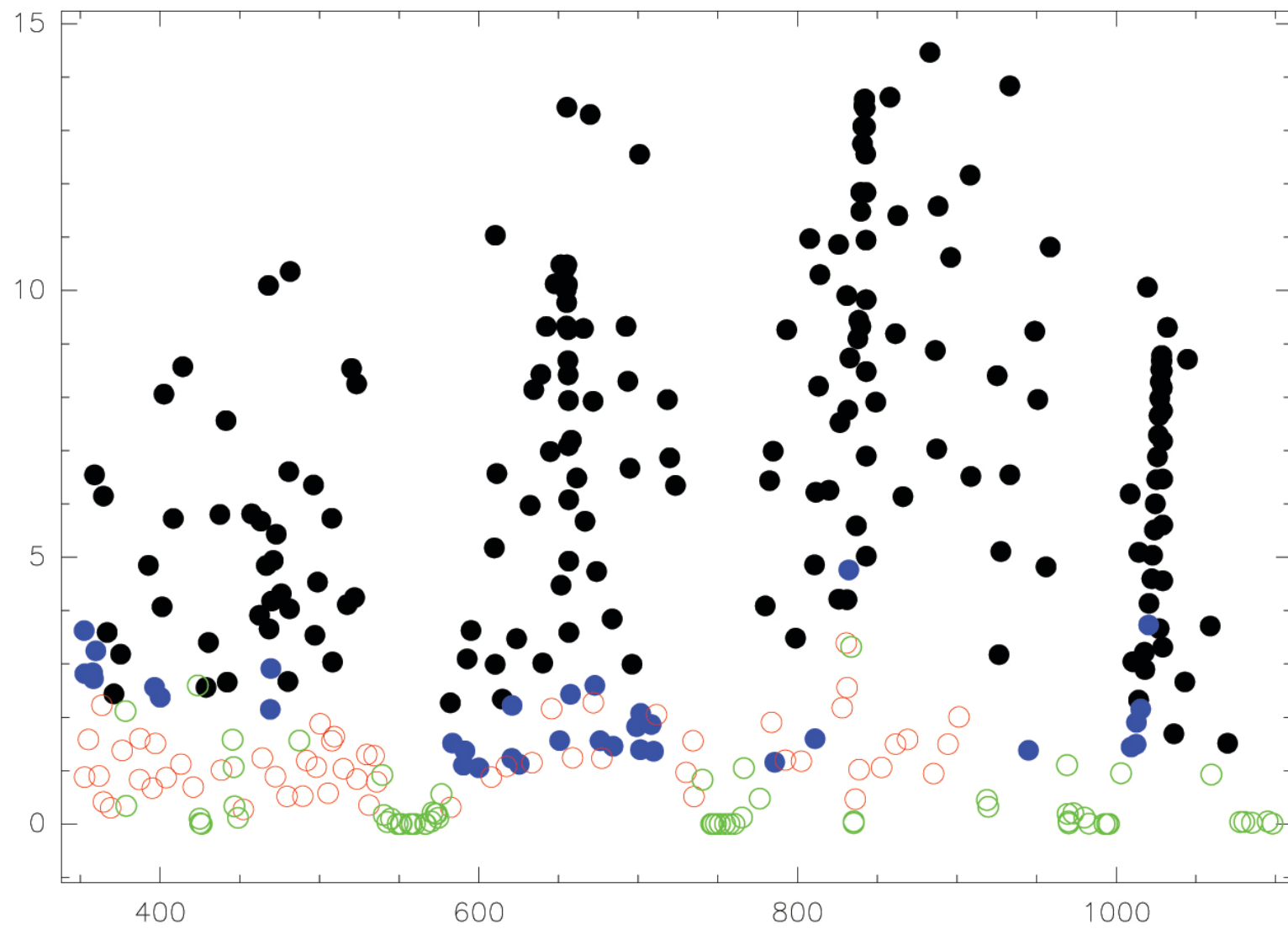


# Relevant results for continuum-like atmospheric absorption below 1 THz

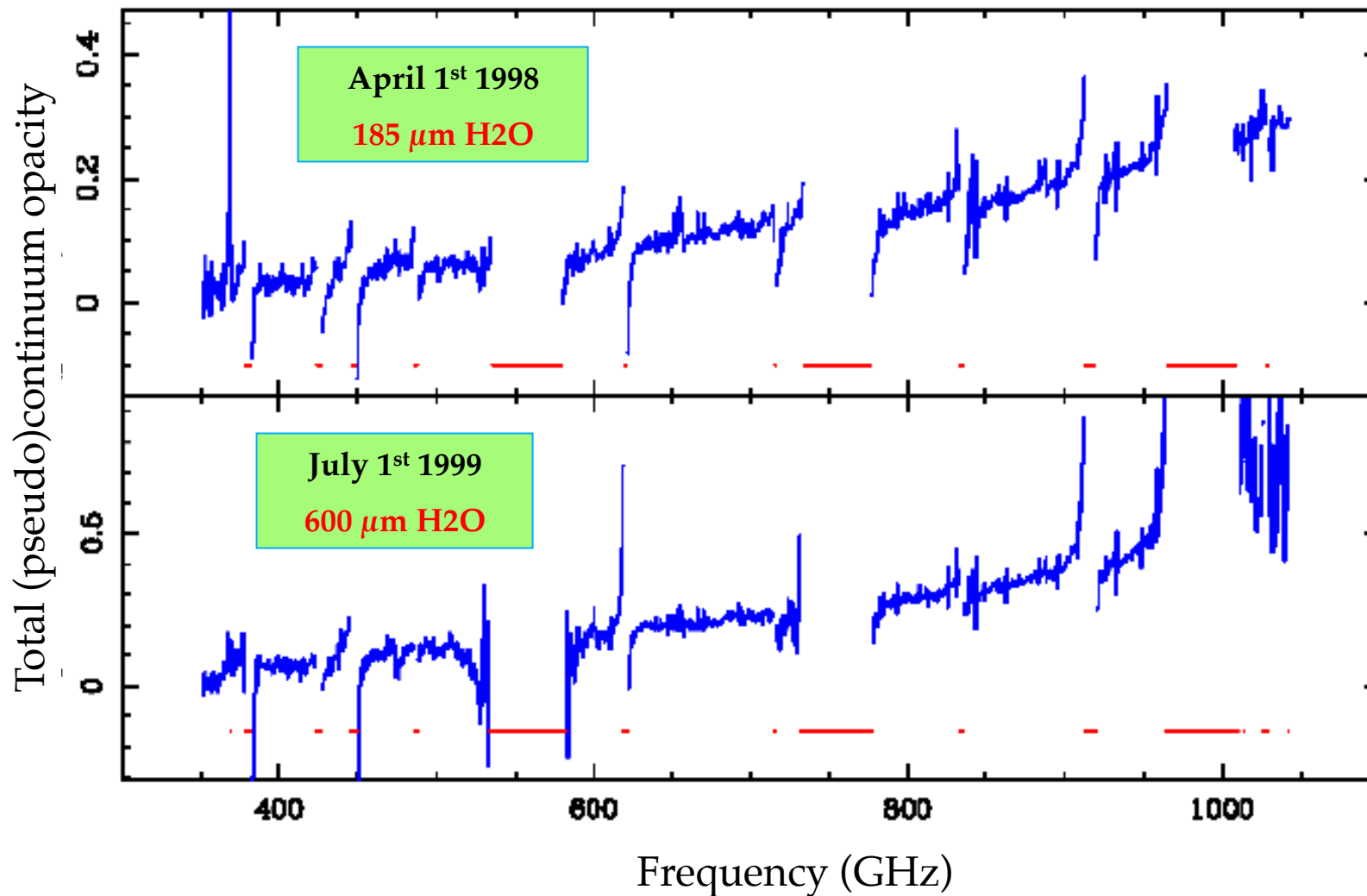


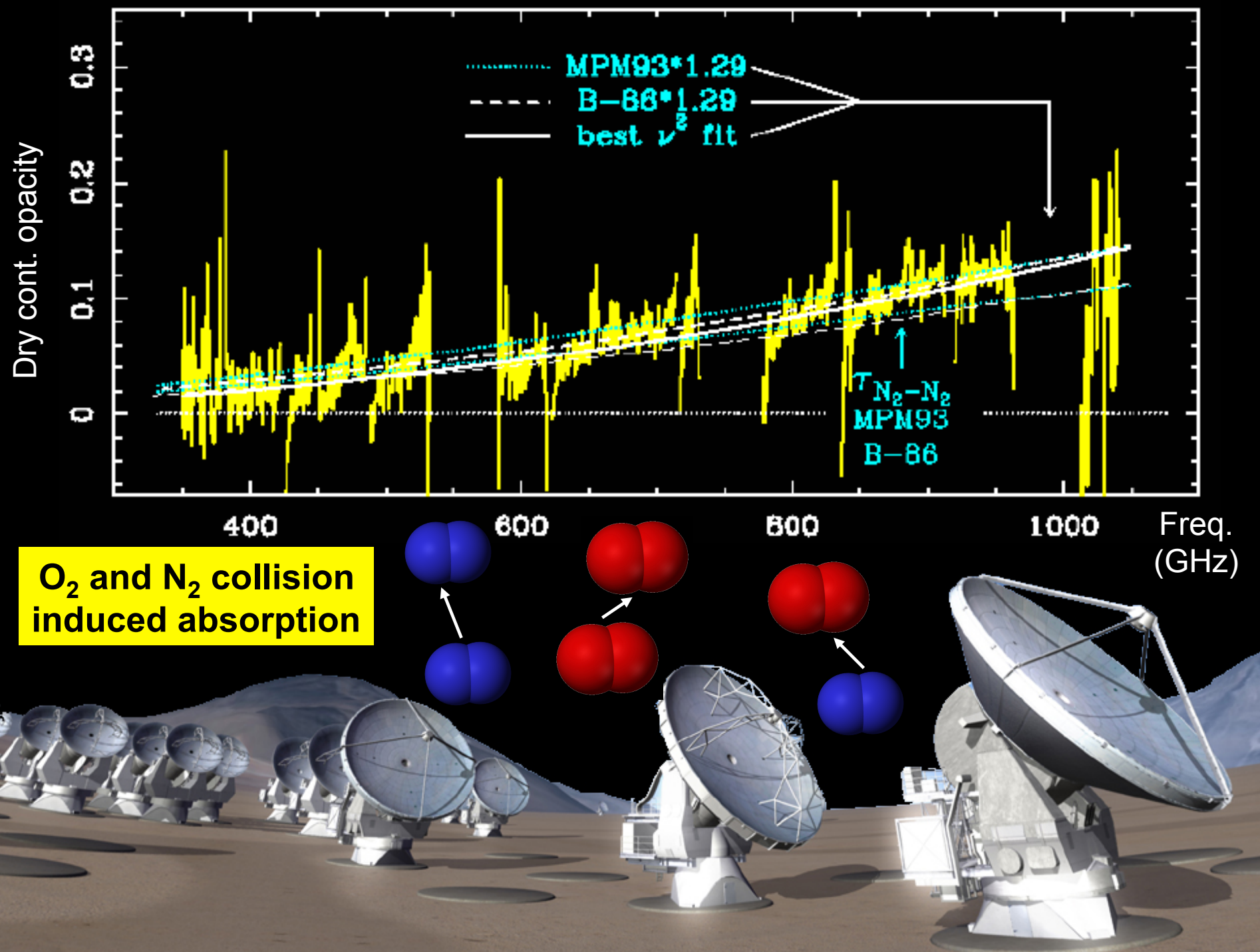




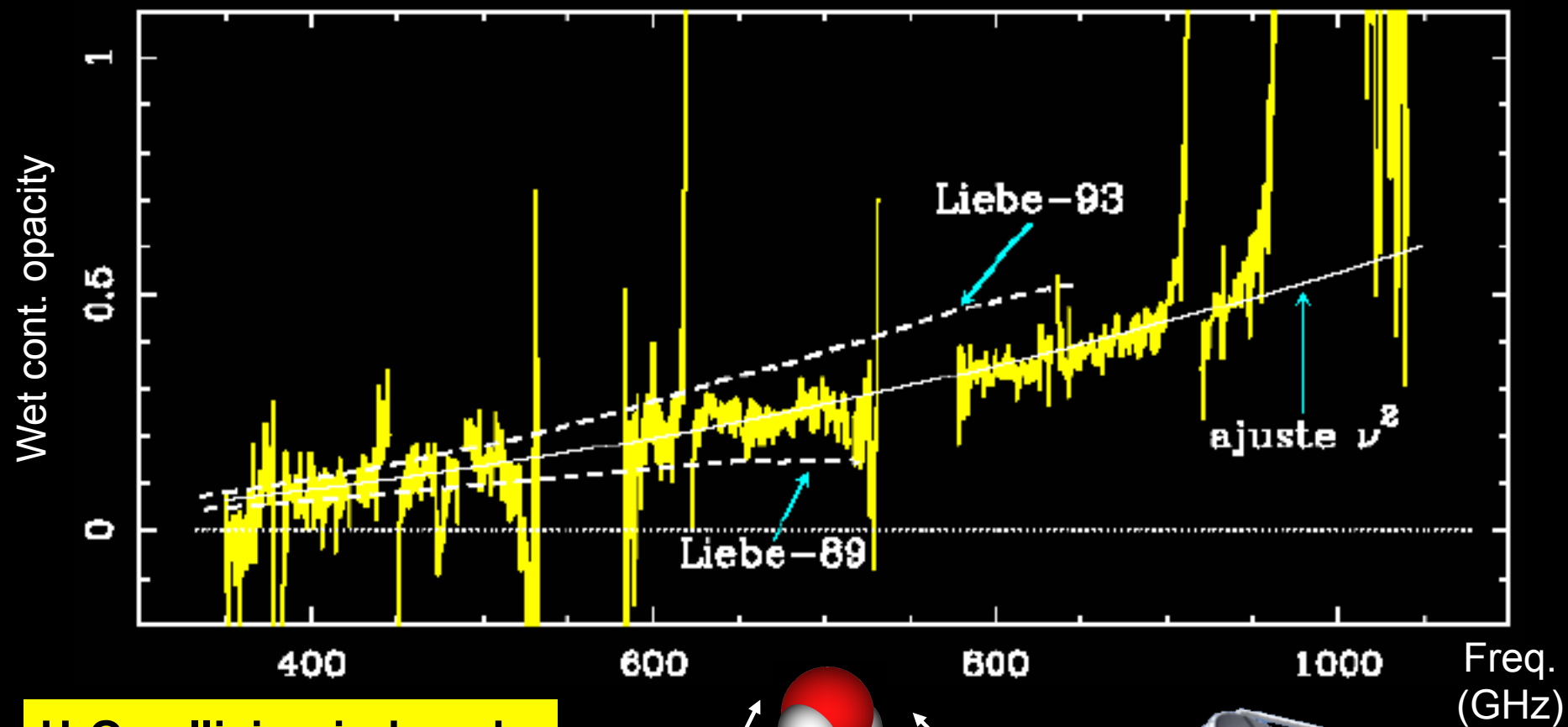


# Continuum-like terms measured with the FTS

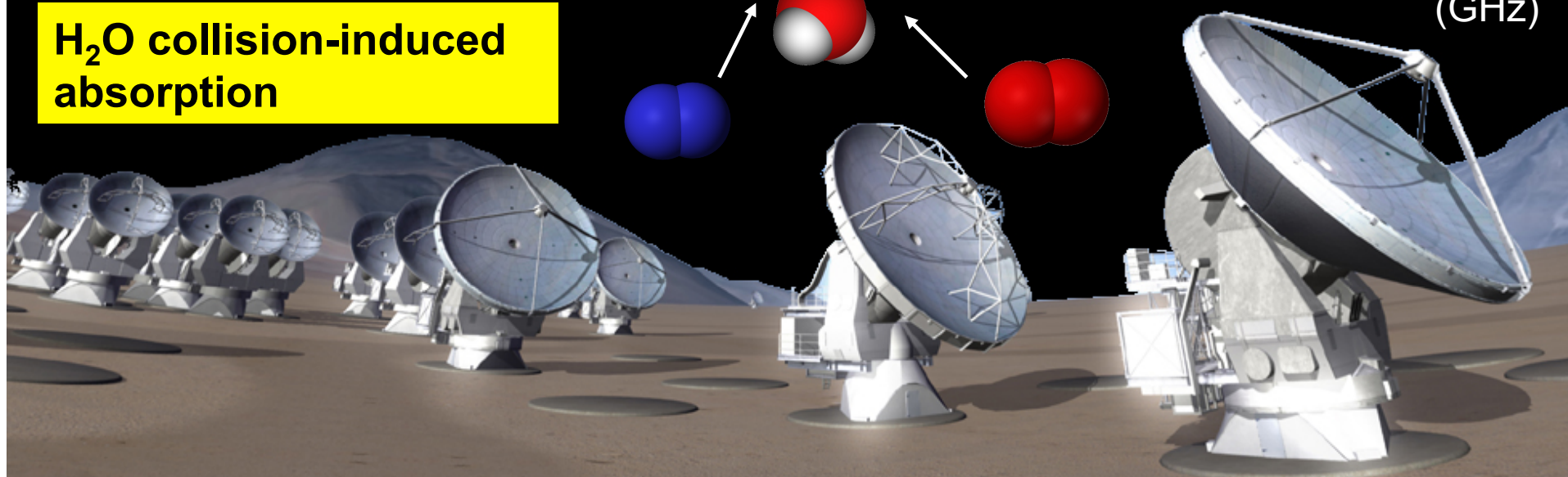
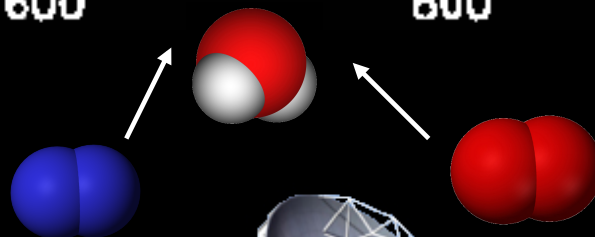








**H<sub>2</sub>O collision-induced absorption**



# **CSO-FTS approach to solve the « excess of continuum » problem**

**Use well calibrated measurements acquired  
during very dry conditions**

- $T_0 \sim 270 \pm 3 \text{ K}$
- $P \sim 620 \pm 1.5 \text{ mb}$  at Mauna Kea summit.

**As a consequence:** The “dry” atmospheric absorption is basically the same (within 1-3 %) in the different situations.  
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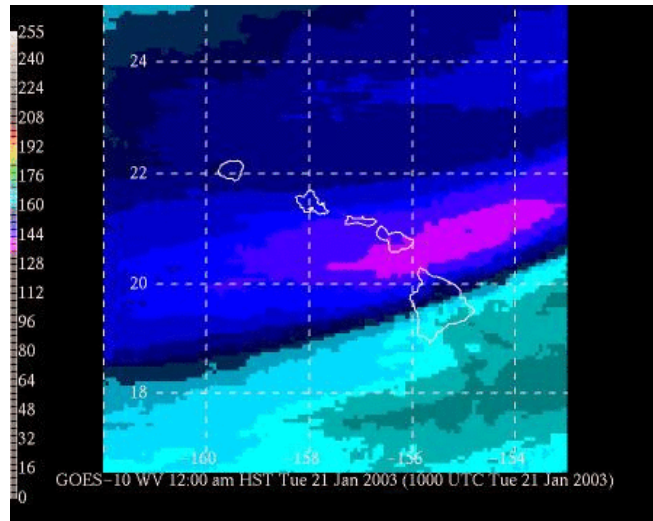
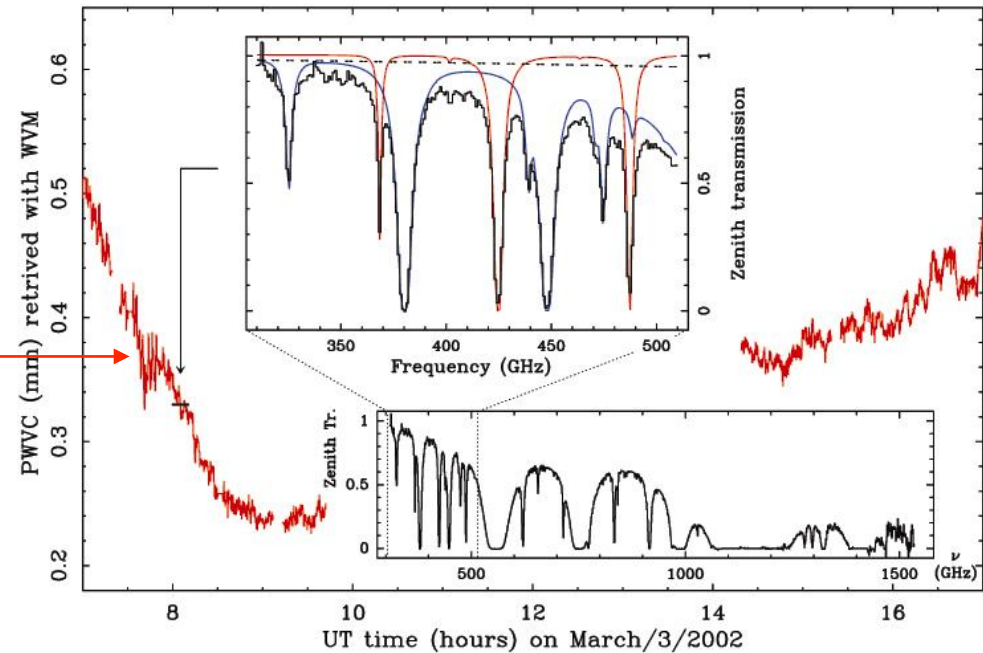
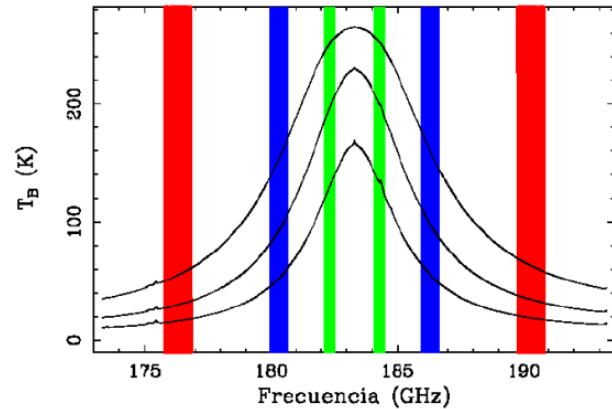
The water vapor column can be determined very precisely from the near wings of water lines, virtually independently of the continuum terms.

# Complementary measurements

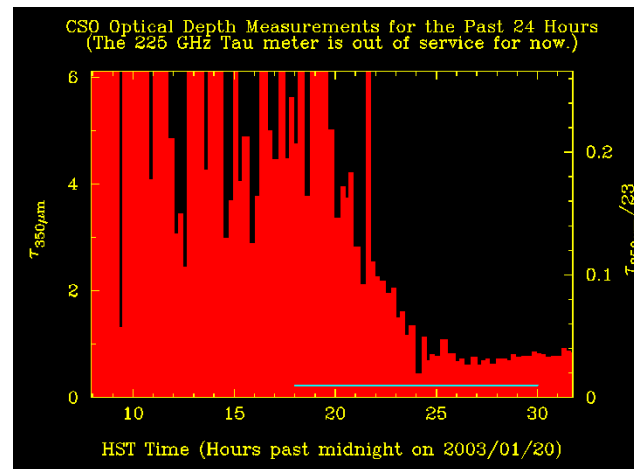
WVM data + FTS data below 500 GHz + GOES-10 data + Tau\_meter data + weather station data: PWV can be determined independently of uncertainties in continuum-like terms and cross-checked

## 183 GHz WV Monitor - Martina Wiedner

- 3 channels, uncooled  $T_{\text{sys}}=2000\text{-}2500\text{K}$ , mounted on telescope, calibration on 300 K and 370 K at 1Hz



GOES-10 Water vapor

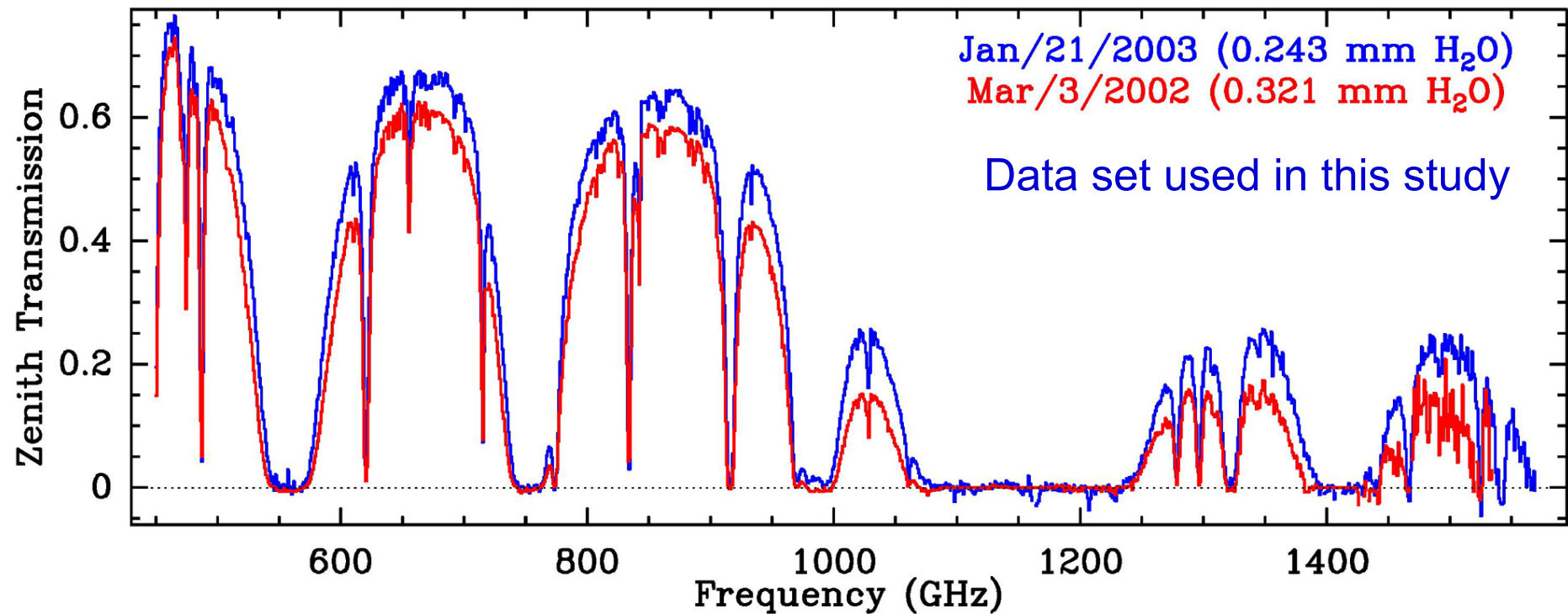


350  $\mu\text{m}$  opacity meter

OTHER:

Hand-held thermo-hygrometer

Telescope's weather station



**WVM data + FTS data below 500 GHz:**

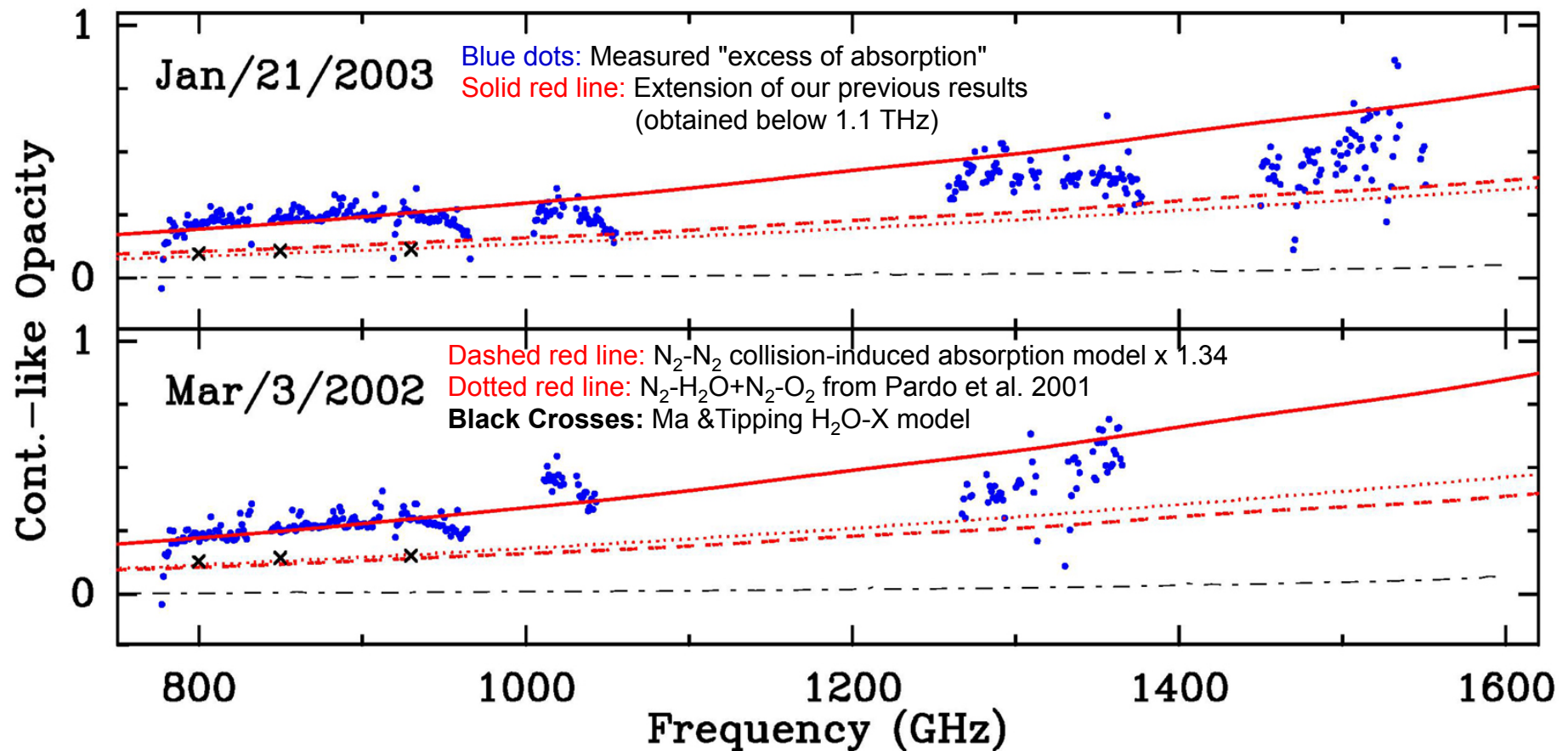
**Allow a Water Vapor Column determination independent of  
uncertainties in continuum-like terms**

**GOES-10 data + Tau\_meter data + weather station data**

**Used as extra information to check consistency**

**Line opacity can be removed considering calculated and/or  
laboratory line parameters. The remaining absorption is the  
"excess of opacity"**





## ¿What is the nature of this opacity term?

- Errors in far wings of lines above 2 THz can be ruled out
- Earlier (Pardo et al 2001) collision-induced absorption explanation confirmed below 1.1 THz
- Extension of those results above 1.1 THz results on overestimation
- **IF**  $\text{N}_2\text{-N}_2 \times 1.34$  collision induced absorption model (Boissoles et al. 2003) is assumed correct **THEN** the  $\text{H}_2\text{O-O}_2 + \text{H}_2\text{O-N}_2$  collision induced absorption calculations (Ma & Tipping 2002) provide very good results below 1 THz and it appears that above that frequency we may start seeing the flattening of this term.

# GENERAL CONCLUSIONS

- An accurate determination of **WVC** from **FTS** or **WVM** data can be done only if the continuum opacity is well known or unimportant (below 0.5 GHz is best).
- Below 1.1 THz the continuum-like opacity has been measured and successfully separated into wet and dry parts in Pardo et al. 2001.
- Data presented here (both **FTS** and **WVM**) have allowed to separate the total continuum from the lines in the range 1.0-1.6 THz.

Serabyn, Weisstein, Lis & Pardo, Appl. Optics, 37:2185, 1998

Matsushita, Matsuo, Pardo, & Radford, PASJ, 51:603, 1999.

Pardo, Serabyn & Cernicharo, JQSRT, 68:419, 2001

Pardo, Cernicharo & Serabyn, IEEE TAP, 49:1683, 2001

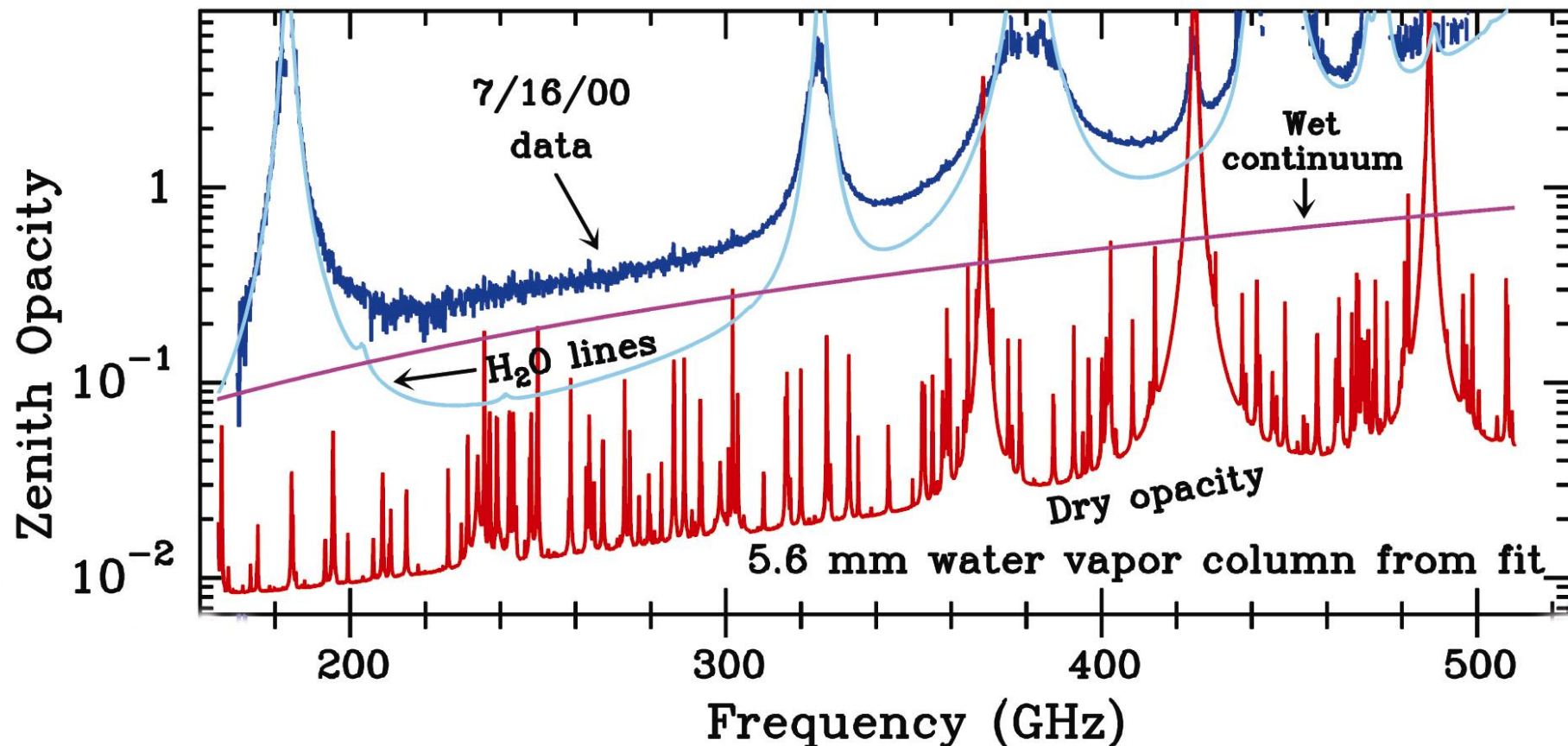
Pardo, Wiedner, Serabyn, Wilson, Cunningham, Hills & Cernicharo, Ap. J. Suppl., 153:363, 2004

Pardo, Serabyn, Wiedner & Cernicharo, JQSRT, 96:537, 2005

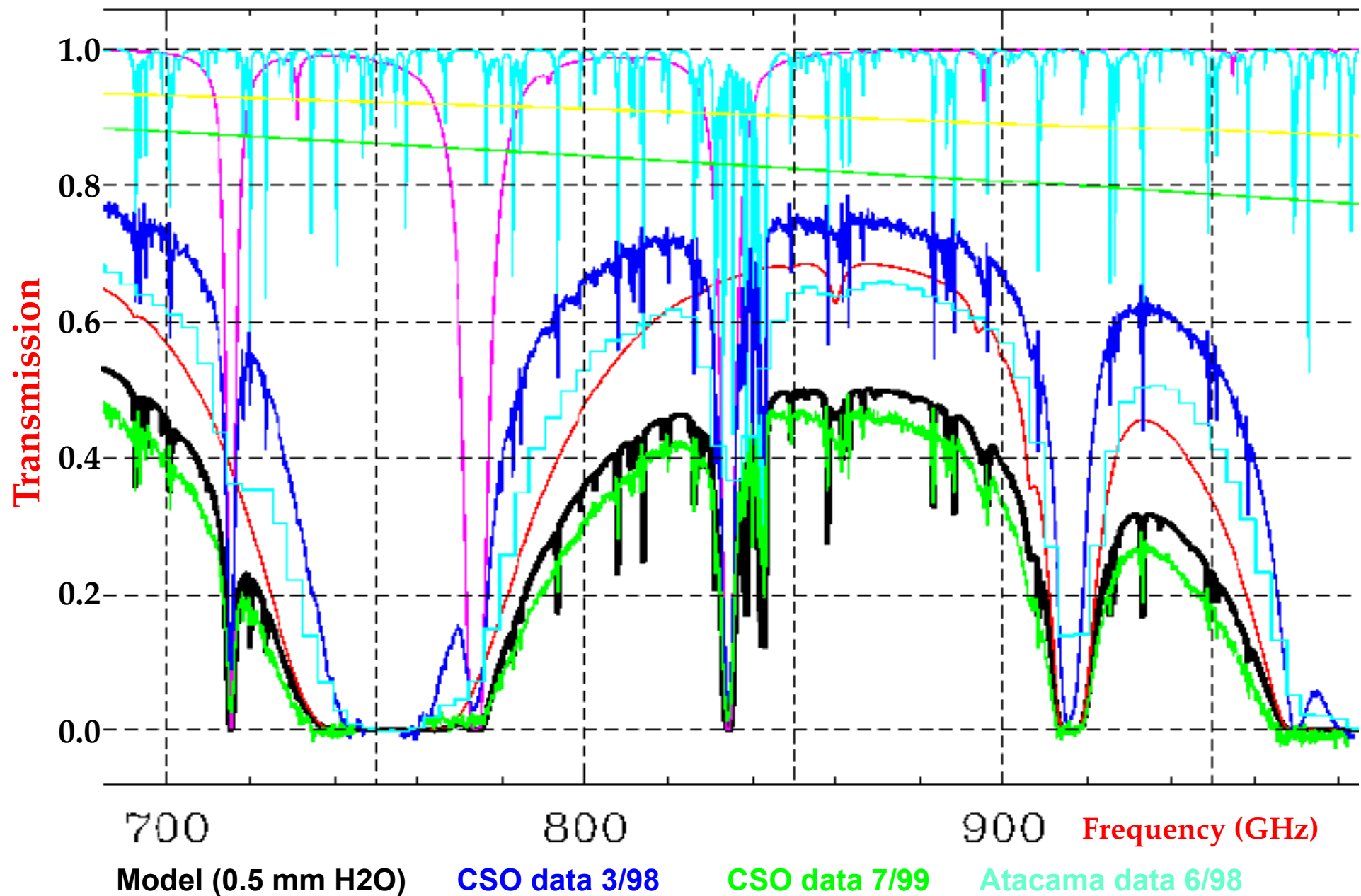
## 1c. Up-to-date model: ATM



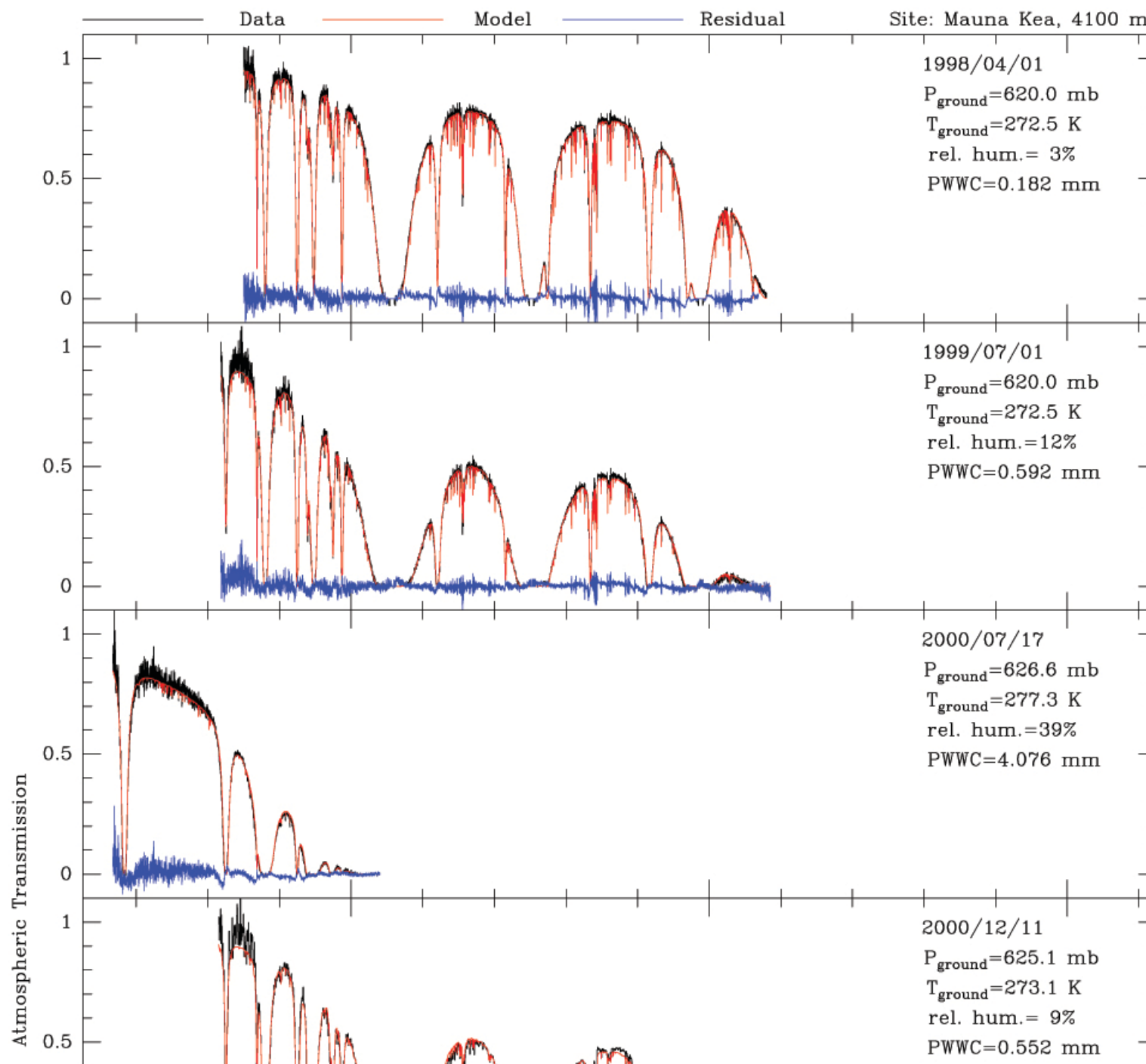
### Example 1: Opacity terms in wet atmospheric conditions at Mauna Kea

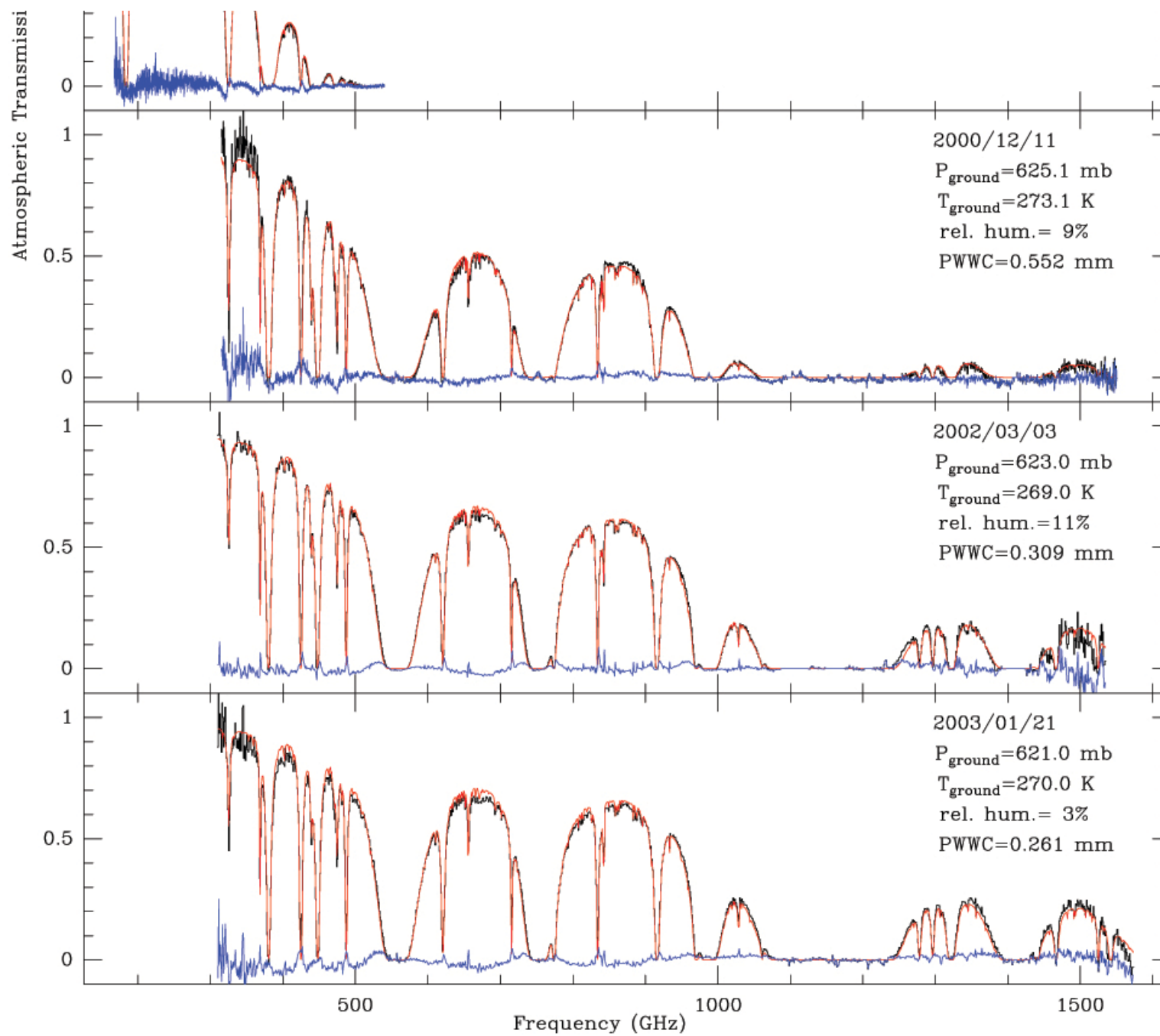


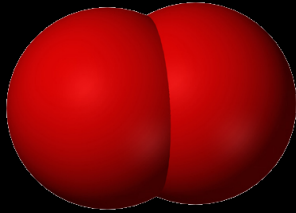
## Example 2: Atmospheric transmission at high frequencies in dry Mauna Kea conditions







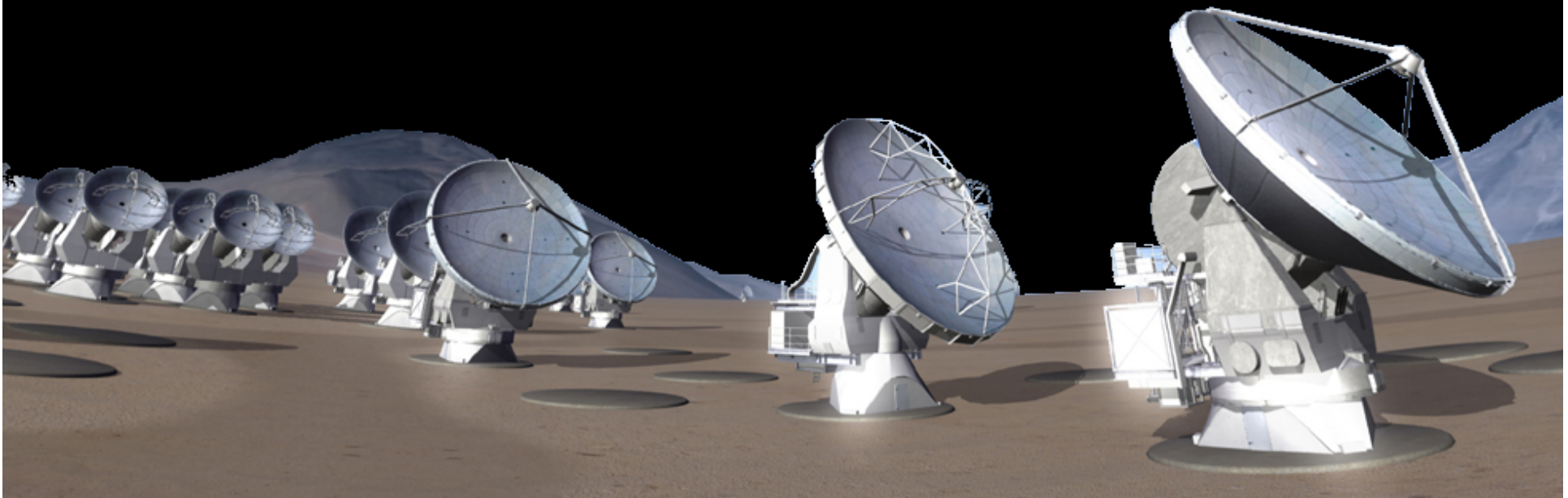




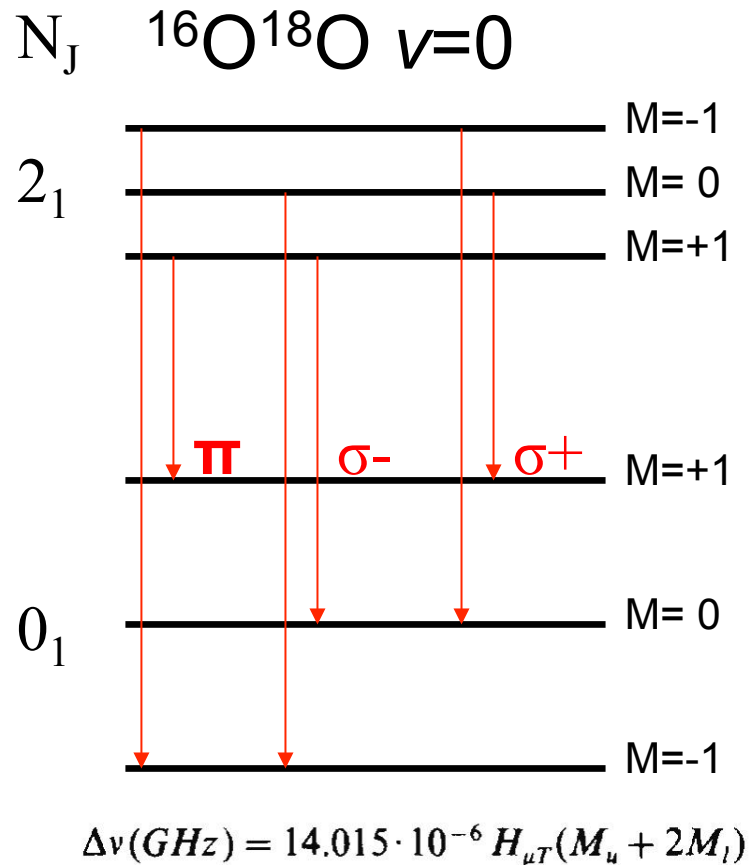
O<sub>2</sub>

Paramagnetic molecule: Coupling of its permanent dipole moment with an external magnetic field causes ZEEMAN SPLITTING.

Modeling this effect is rather complex because of anisotropy, polarization, etc...







**Transitions  $\pi$  ( $\Delta M=0$ ):** Radiation linearly polarized in the direction of the geomagnetic field.

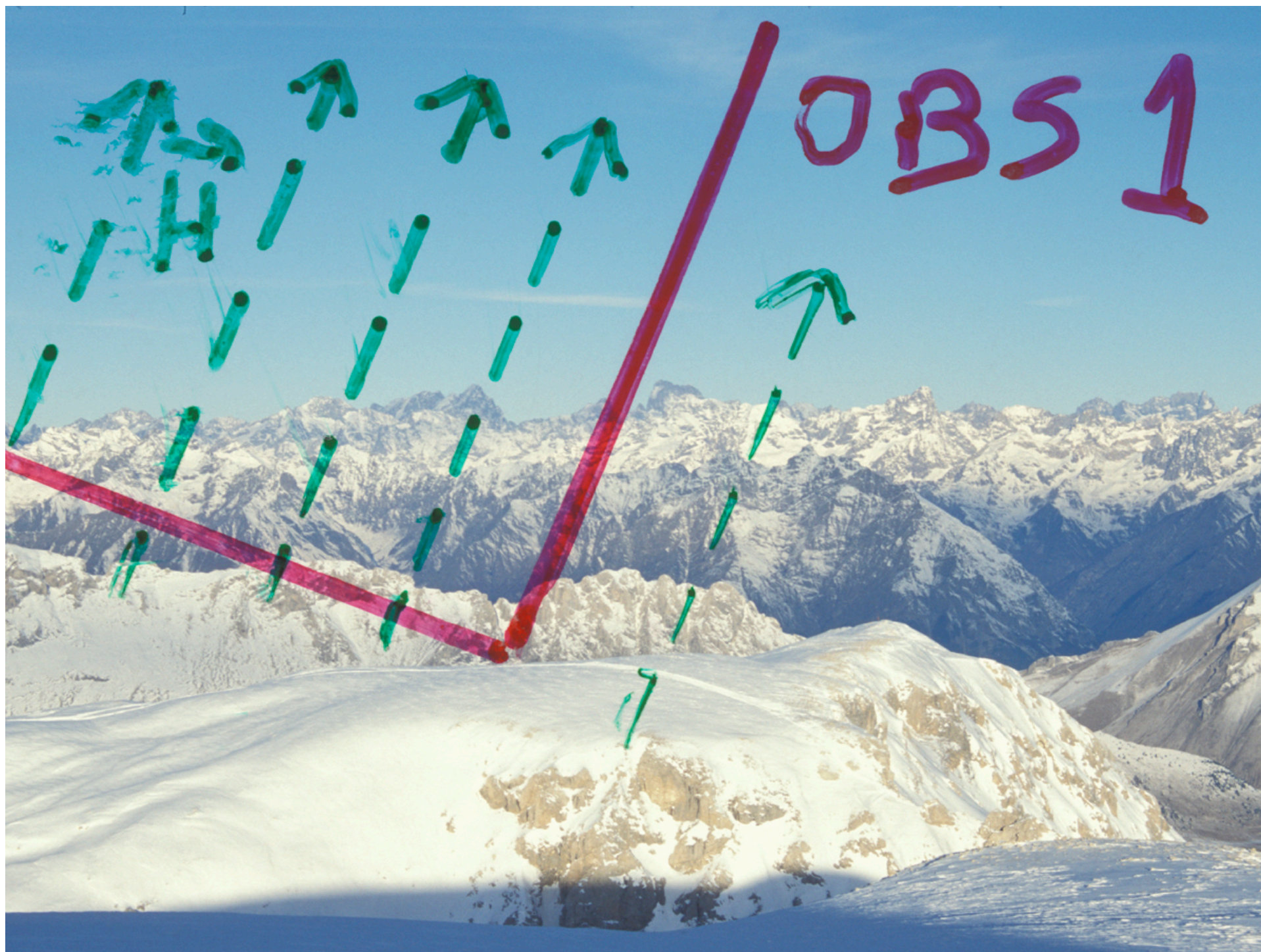
**Transitions  $\sigma$  ( $\Delta M=\pm 1$ ):** Radiation circularly polarized (right-hand or left-hand in the plane perpendicular to the direction of the geomagnetic field).

We should expect differences in the line profile depending on the line of sight, the type of polarization detected, and the orientation of our detector with the geomagnetic field.

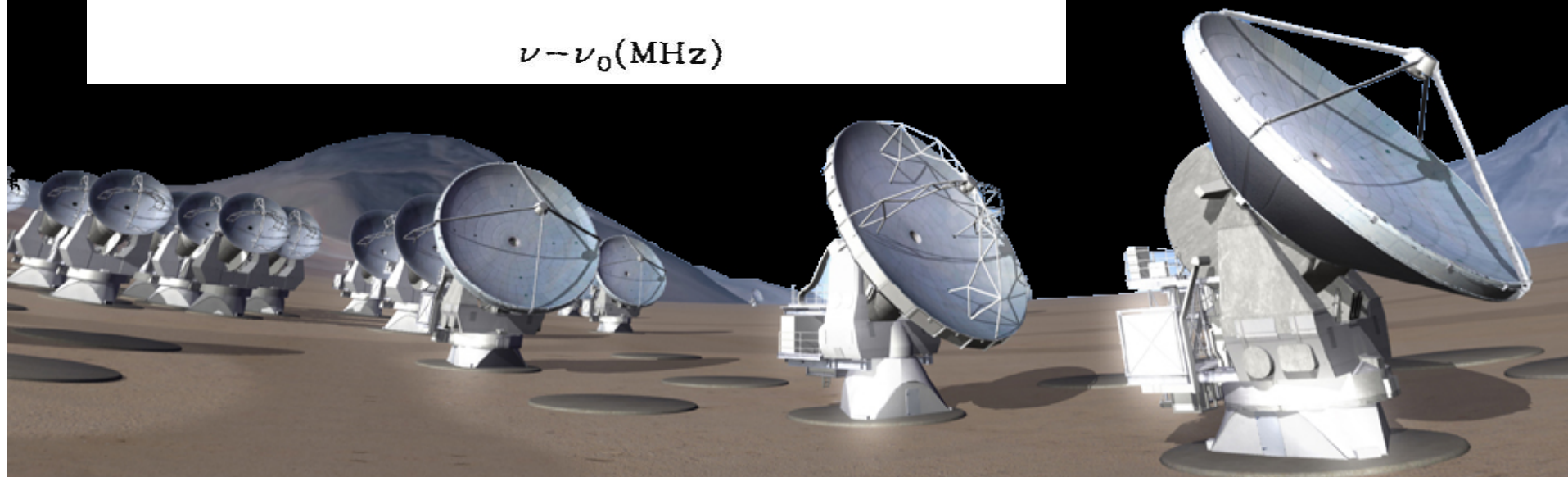
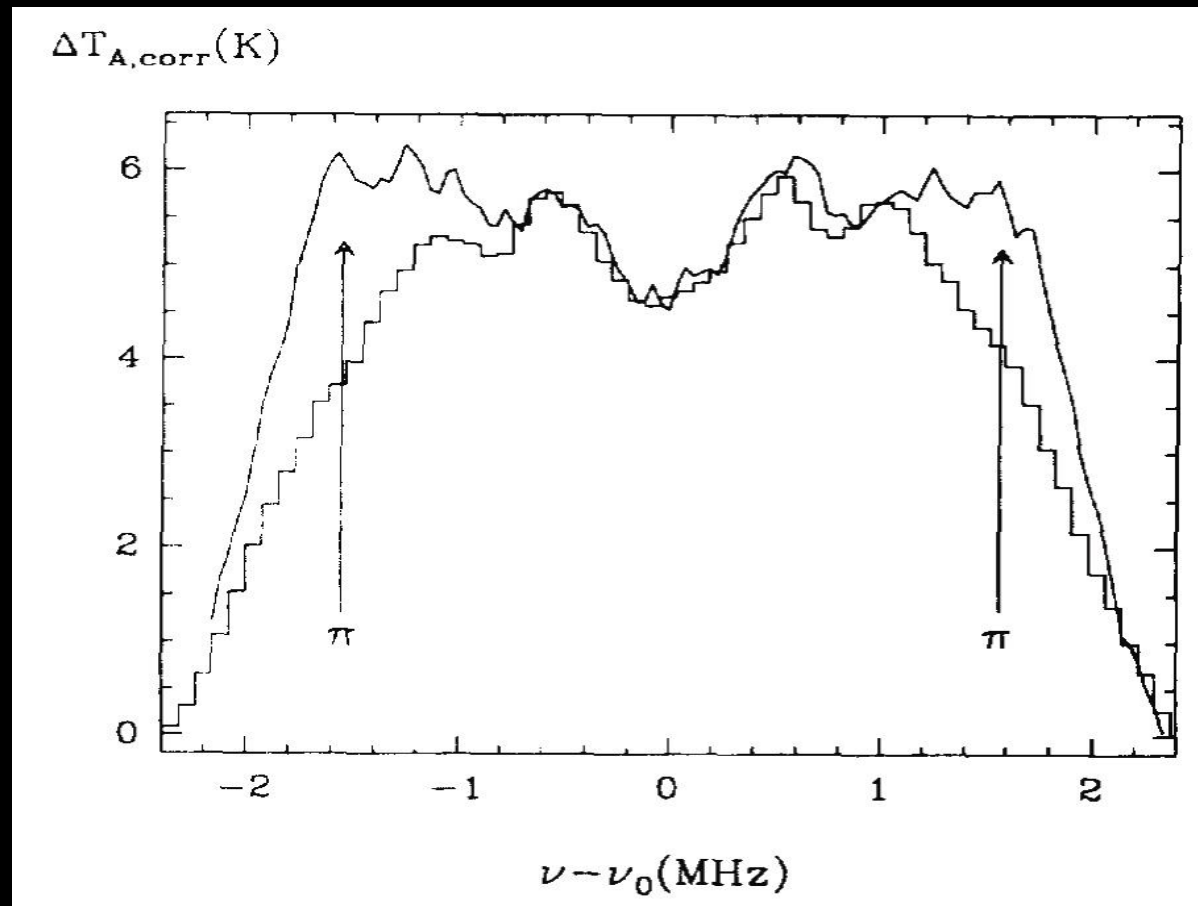


**Experiment at POM2. SIS receiver and autocorrelator providing 39 kHz resolution and 4.53 MHz of total band.**





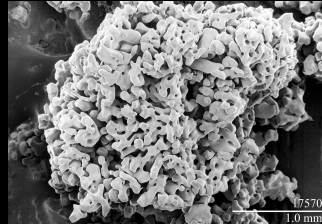




# Hydrometeors contribution (absorption, scattering, phase...)



aerosols



Wet snow



snow

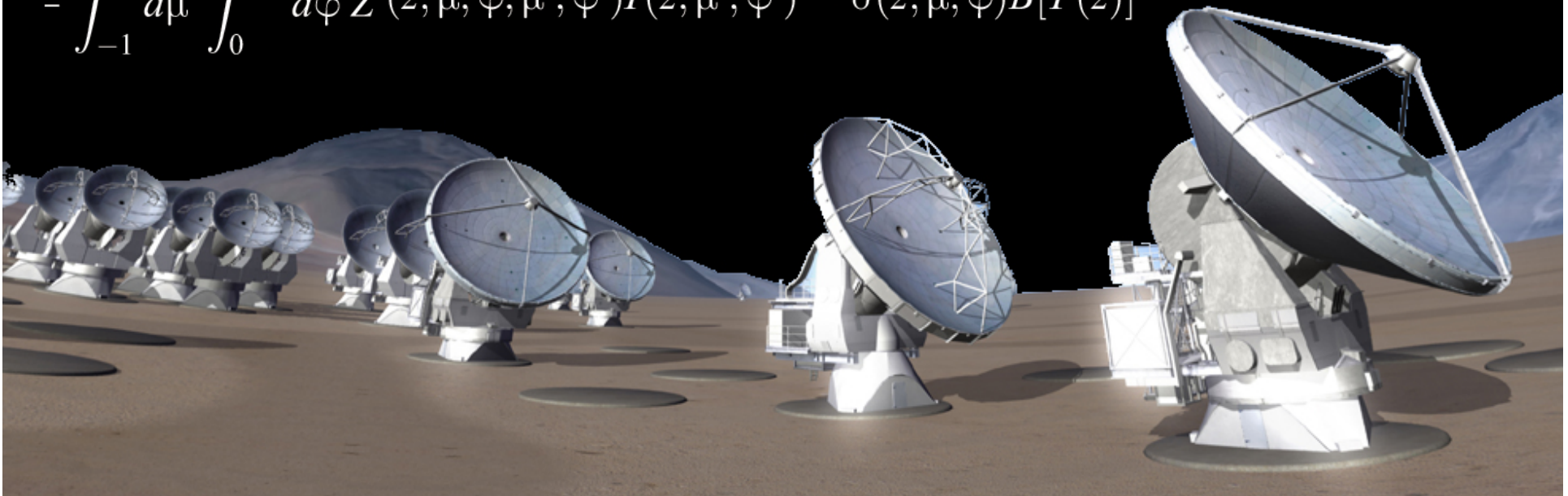


Liquid water



Hale

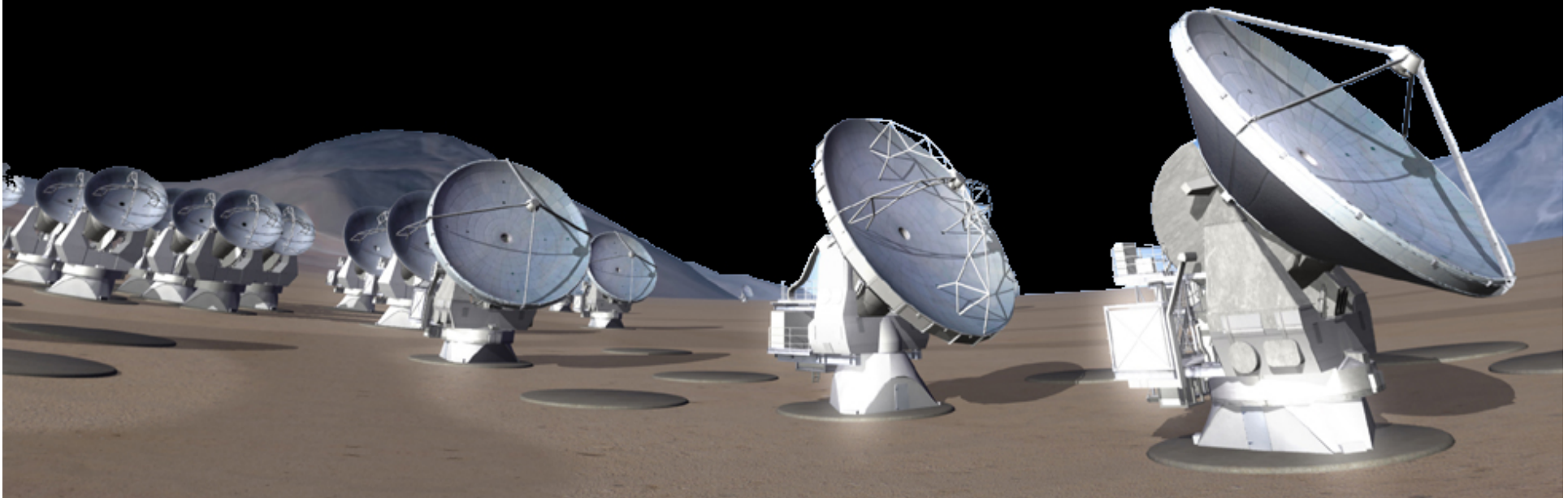
$$\mu \frac{dI(z, \mu, \varphi)}{dz} = K(z, \mu, \varphi)I(z, \mu, \varphi) - \int_{-1}^1 d\mu' \int_0^{2\pi} d\varphi' Z(z, \mu, \varphi, \mu', \varphi') I(z, \mu', \varphi') - \sigma(z, \mu, \varphi) B[T(z)]$$



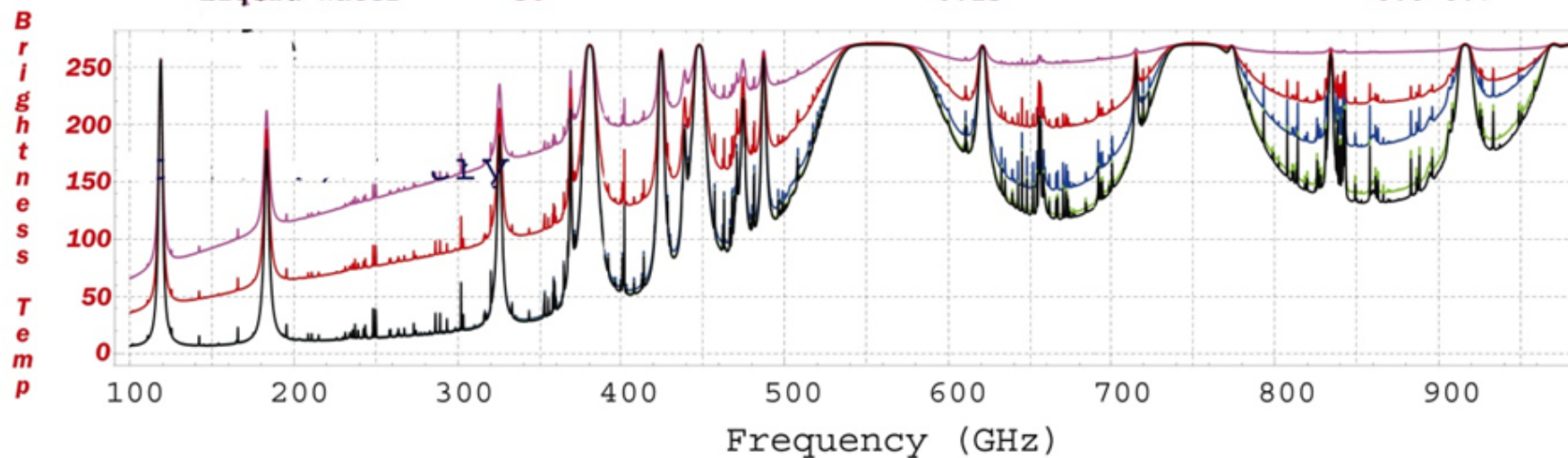


## In ATM:

- Phase Matrix calculations from M. I Mishchenko (prolate and oblate spheroids with azimuthally random distribution, or T-matrix method for spheres.
- Refraction indexes from literature.
- RT using DDA method from Evans et al.
- Single scattering assumed within each layer.
- Lambertian, Fresnel and other surface types.



Cloud type	Part. radius (microns)	Equiv. Water Path(mm)	Cloud Position (km)
Clear Atm.	----	----	----
Ice	30	0.07	6.0-6.7
Ice	50	0.15	6.0-6.7
Liquid Water	30	0.07	6.0-6.7
Liquid Water	50	0.15	6.0-6.7



# Atmospheric Phase fluctuations



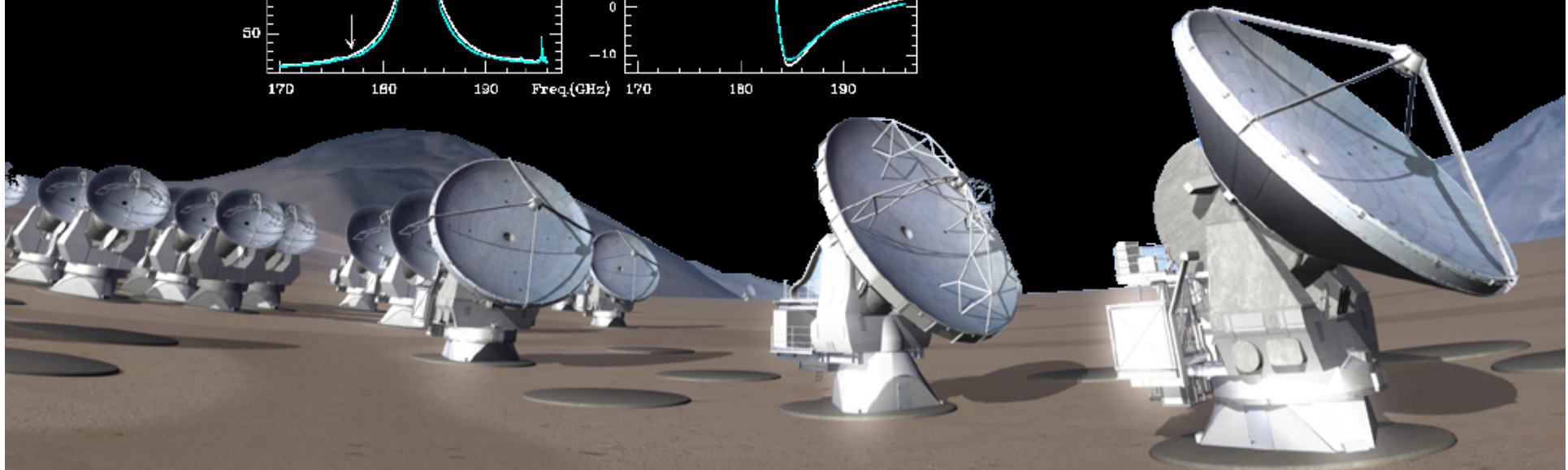
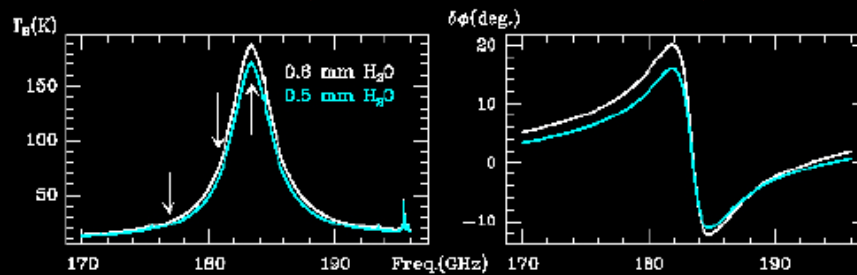
$$(\kappa_\nu)_{lu} = \frac{8\pi^3 N \nu}{3hcQ} \left( e^{-E_l/kT} - e^{-E_u/kT} \right) \cdot |\langle u | \mu | l \rangle|^2 f(\nu, \nu_{l \rightarrow u})$$

$\kappa_\nu$  Is a complex number

$$\mathcal{F}(\nu, \nu_{u \leftrightarrow l}) = \frac{\nu}{\pi \nu_{u \leftrightarrow l}} \left[ \frac{1 - i\delta}{\nu_{u \leftrightarrow l} - \nu - i\Delta\nu} + \frac{1 + i\delta}{\nu_{u \leftrightarrow l} + \nu + \Delta\nu} \right] \quad (1)$$

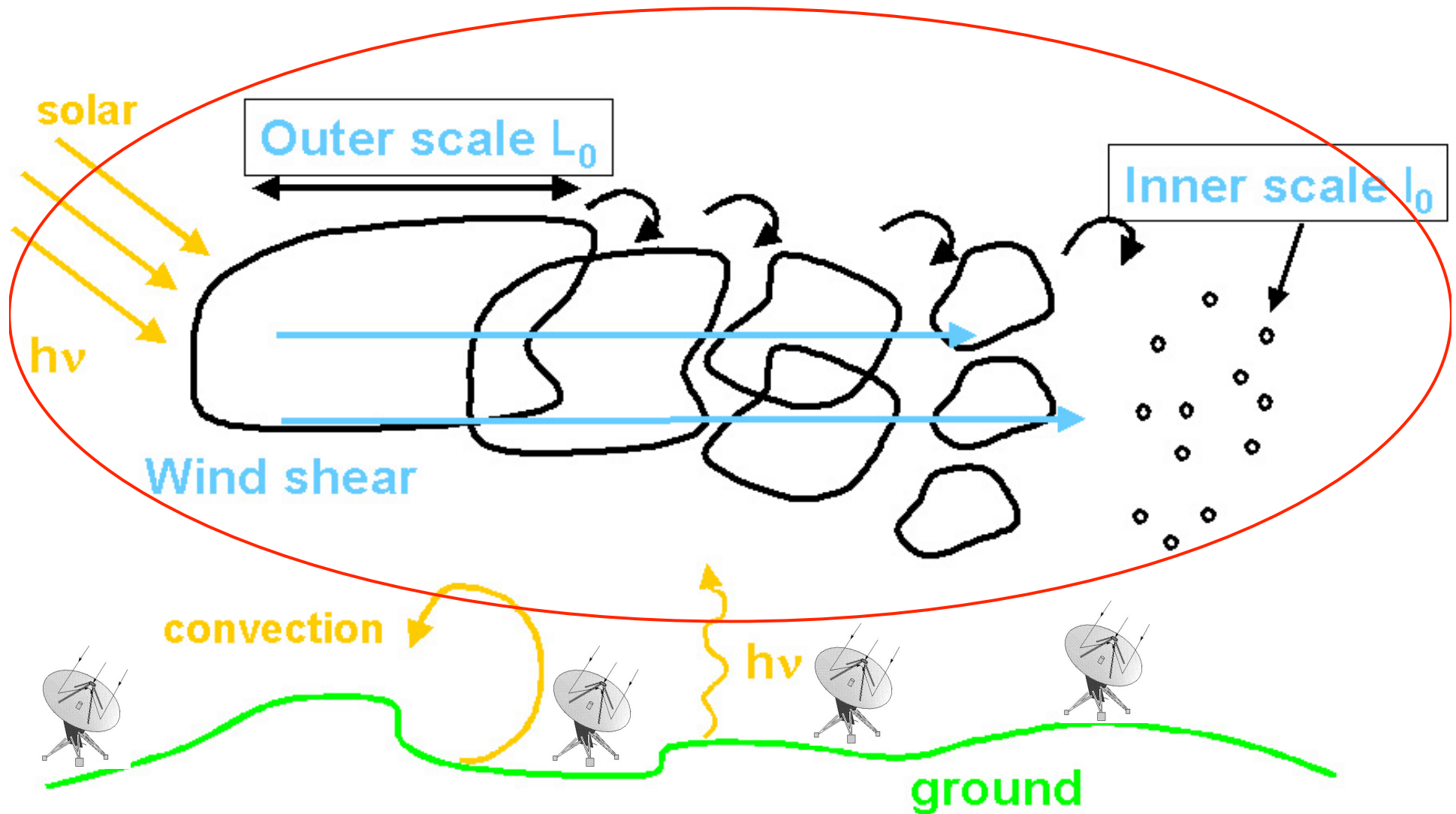
Imaginary part (absorption)

Real part (phase delay or pathlength variation)

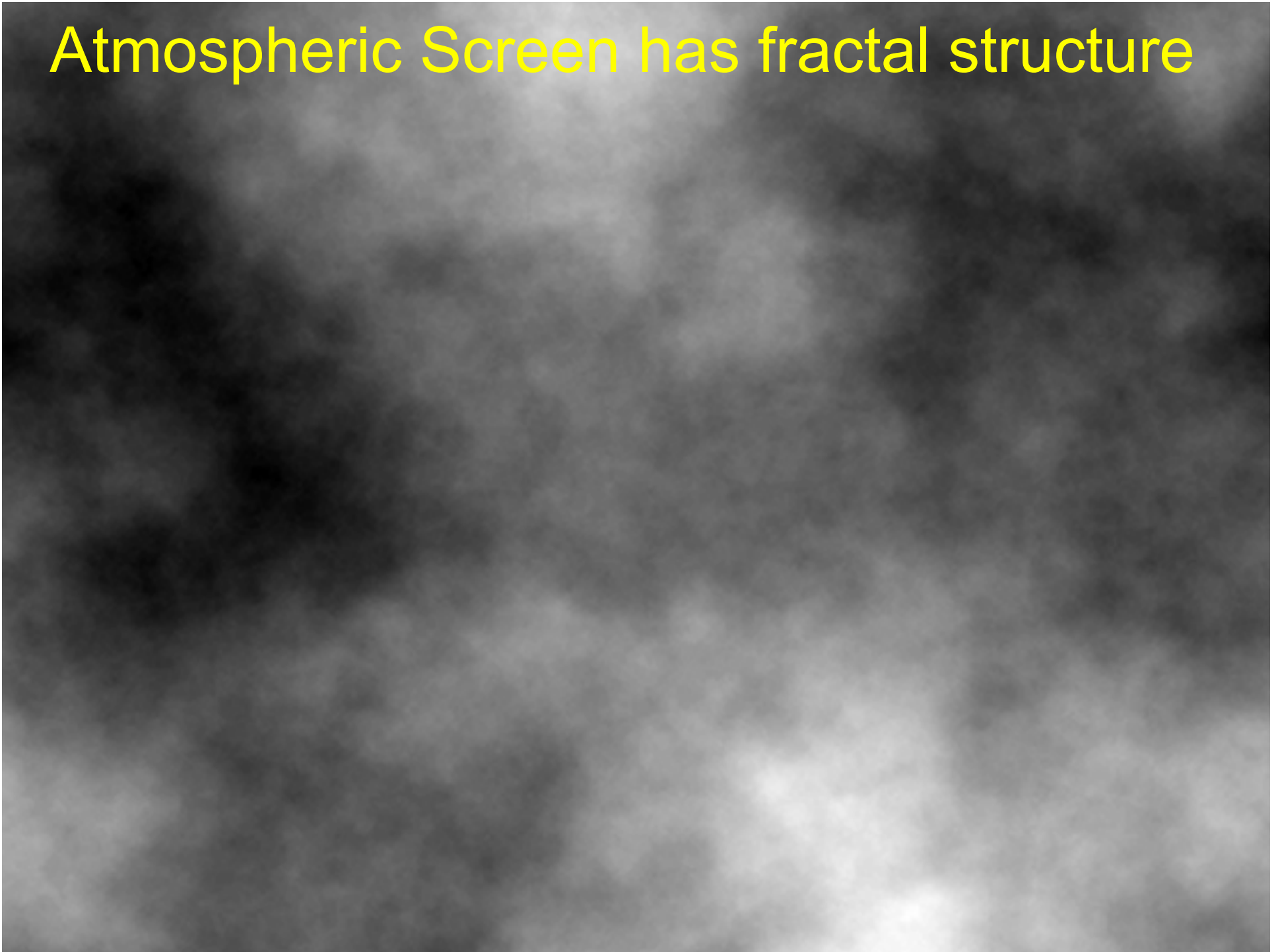




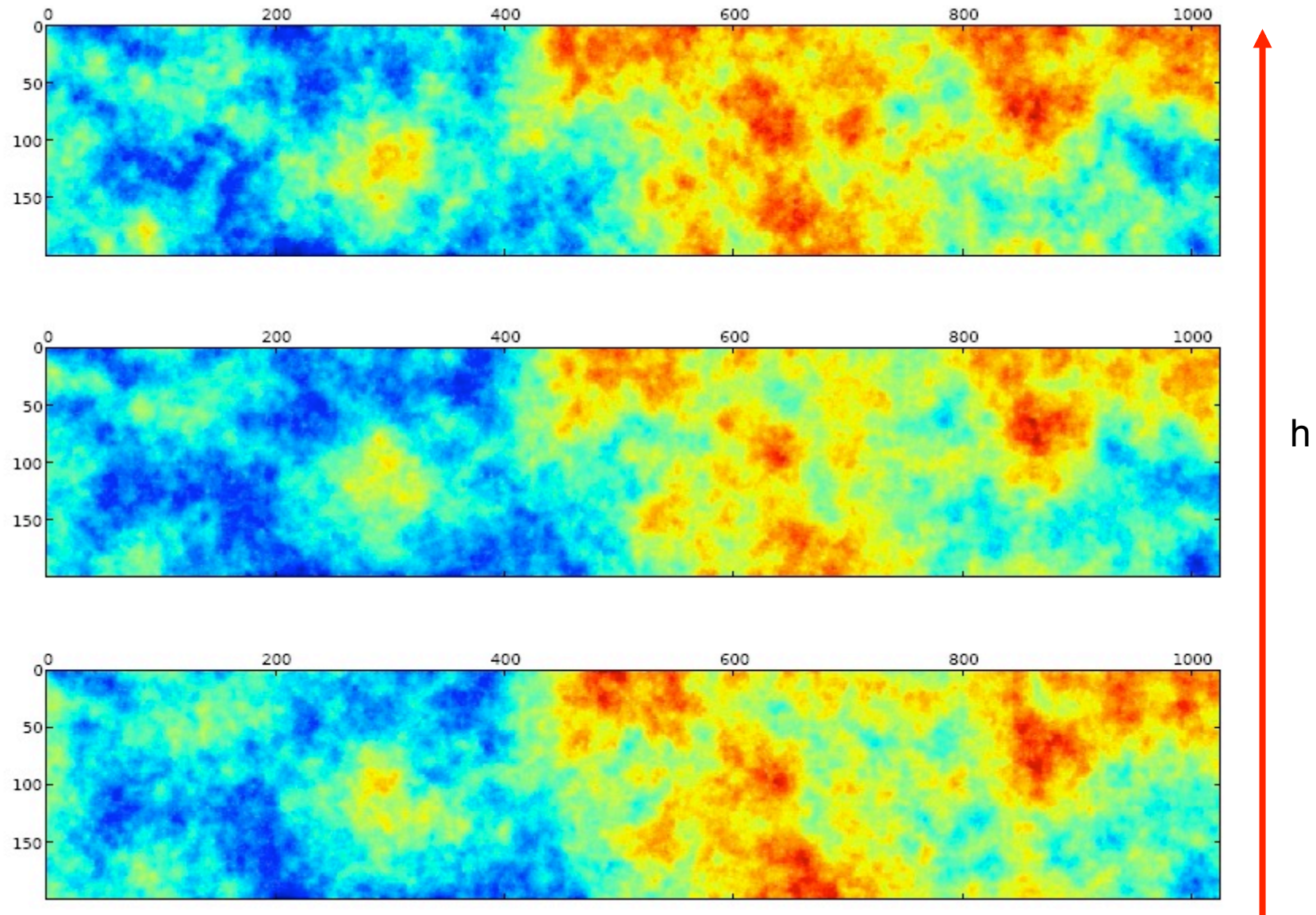
Atmospheric Screen (Wet component moves and evolves in relatively short timescales,  $\sim 1$  sec)



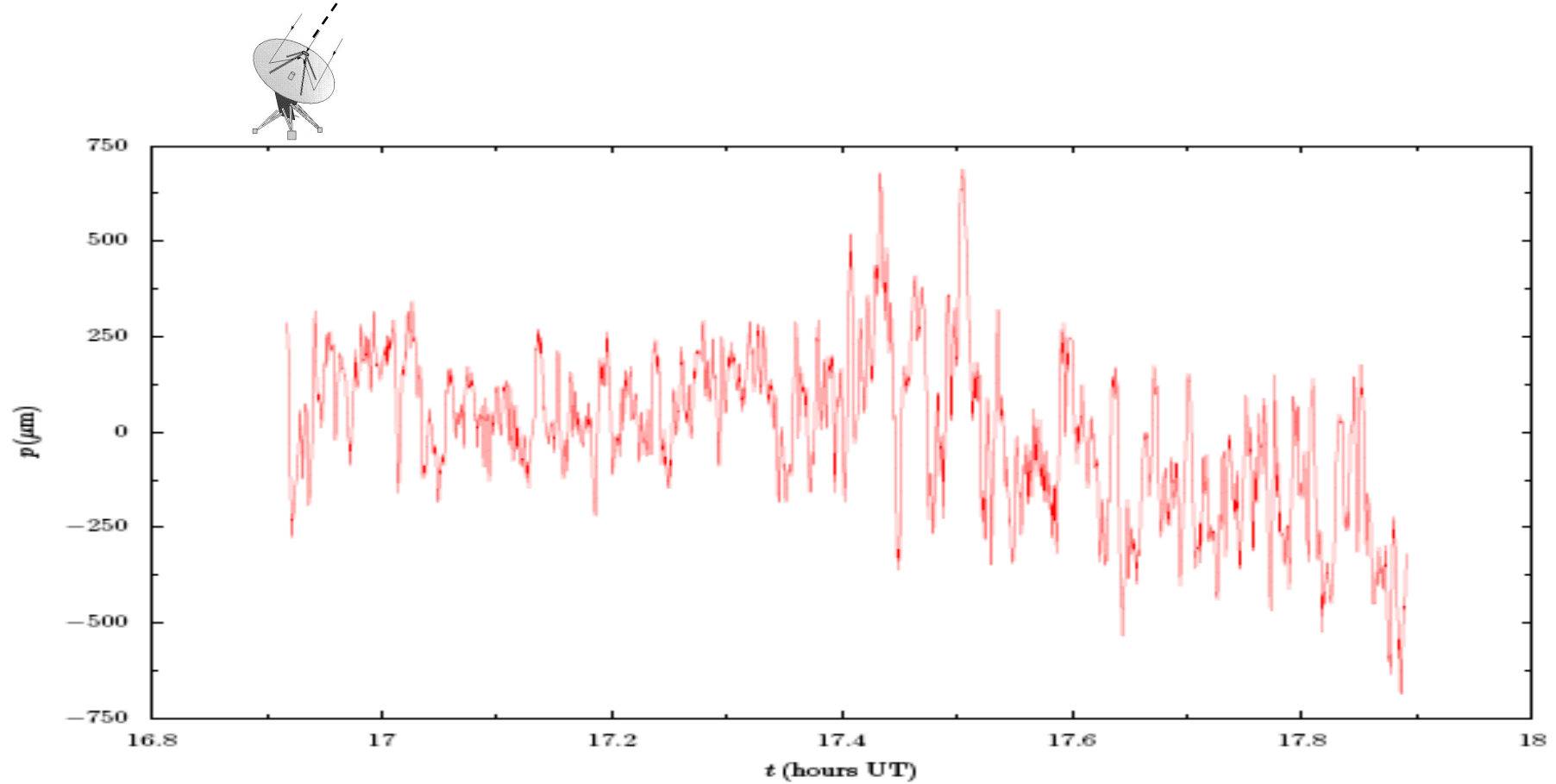
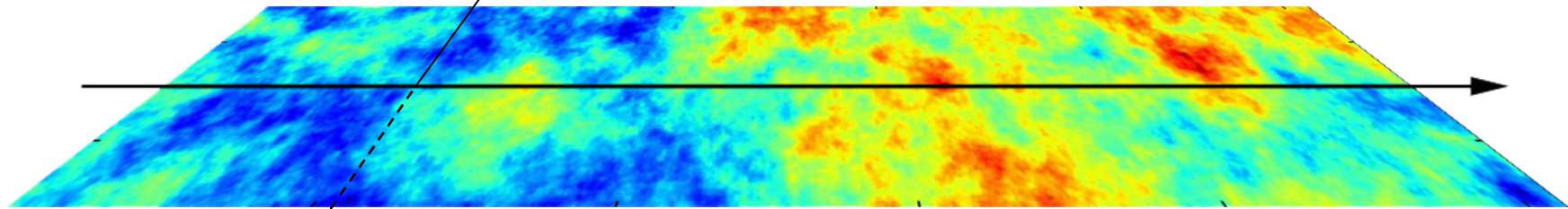
Atmospheric Screen has fractal structure



# Atmospheric screen is three-dimensional



☀ Pathlength fluctuations observed at Mauna Kea while tracking a quasar for 1 hr

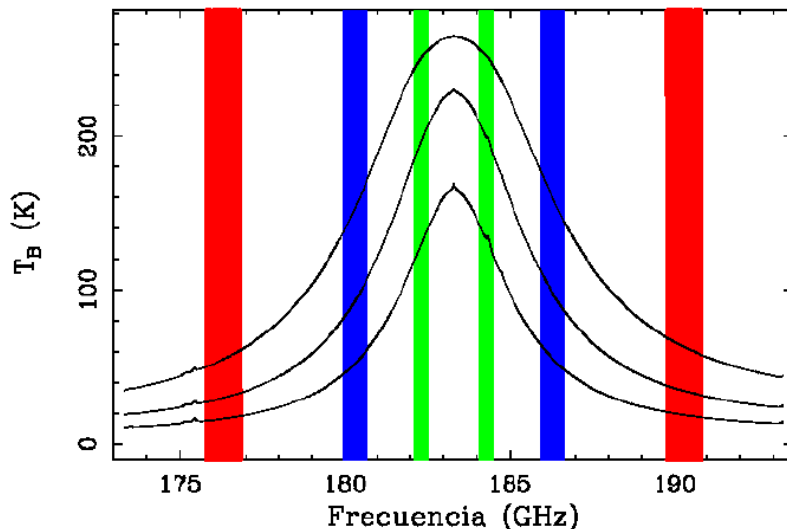


Courtesy of B. Nikolic, Cavendish Laboratory, Cambridge



## Phase correction can be performed using water vapor radiometry at frequencies sensitive to H<sub>2</sub>O

Wavelength (microns)	Pathlength fluctuation factor (microns per micron_H <sub>2</sub> O)
612	8.47
482	6.65
312	10.41



Taking an average Precipitable Water Vapor amount of 0.5 mm, we have:

23 mk/μm    60 mk/μm    173 mk/μm

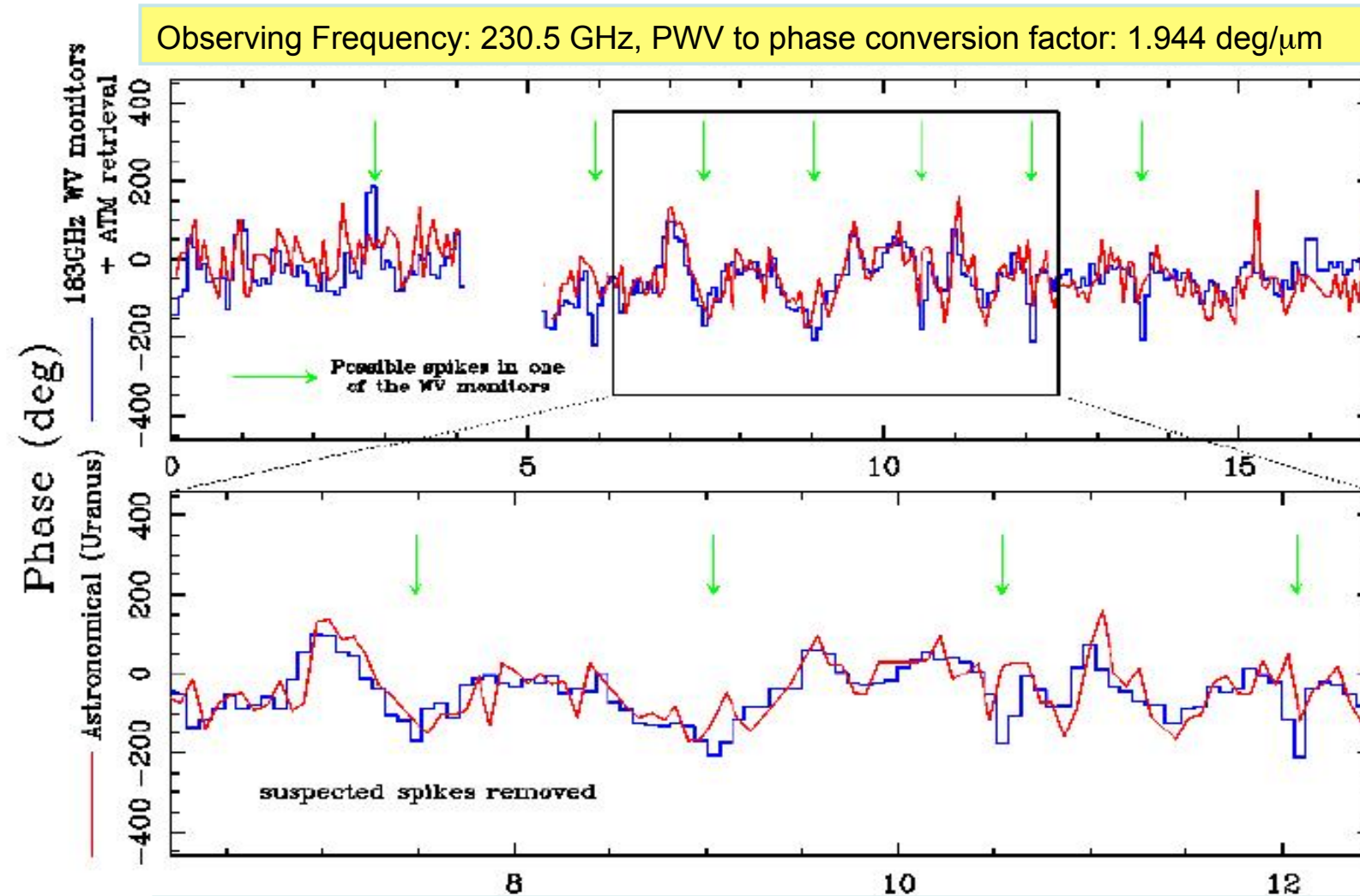
Assuming 0.4 K noise in 1 s in all three channels of the Water Vapor Radiometer

Uncertainty in determining  $\Delta$ PWV: 2.3 μm

The goal of 15 μm pathlength correction accuracy for ALMA should be reachable with this technique in time scales of the order of 1 s.

SYSTEM ALREADY DEVELOPED

## Phase fluctuations: Correction using 183 GHz water vapor radiometers



Time since the beginning of the observation (min) on Nov. 25, 2001



ATM and water vapor radiometry at 183 GHz make ALMA a much powerful observatory at all its frequencies

2016

